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Earth system data cubes unravel global multivariate dynamics

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Abstract. Understanding Earth system dynamics in light of ongoing human intervention and dependency remains a major scientific challenge. The unprecedented availability of data streams describing different facets of the Earth now offers fundamentally new avenues to address this quest. However, several practical hurdles, especially the lack of data interoperability, limit the joint potential of these data streams. Today, many initiatives within and beyond the Earth system sciences are exploring new approaches to overcome these hurdles and meet the growing interdisciplinary need for data-intensive research; using data cubes is one promising avenue. Here, we introduce the concept of Earth system data cubes and how to operate on them in a formal way. The idea is that treating multiple data dimensions, such as spatial, temporal, variable, frequency, and other grids alike, allows effective application of user-defined functions to co-interpret Earth observations and/or model–data integration. An implementation of this concept combines analysis-ready data cubes with a suitable analytic interface. In three case studies, we demonstrate how the concept and its implementation facilitate the execution of complex workflows for research across multiple variables, and spatial and temporal scales: (1) summary statistics for ecosystem and climate dynamics; (2) intrinsic dimensionality analysis on multiple timescales; and (3) model–data integration. We discuss the emerging perspectives for investigating global interacting and coupled phenomena in observed or simulated data. In particular, we see many emerging perspectives of this approach.
for interpreting large-scale model ensembles. The latest developments in machine learning, causal inference, and model–data integration can be seamlessly implemented in the proposed framework, supporting rapid progress in data-intensive research across disciplinary boundaries.

1 Introduction

Predicting the Earth system’s future trajectory given ongoing human intervention into the climate system and land surface transformations requires a deep understanding of its functioning (Schellnhuber, 1999; IPCC, 2013). In particular, it requires unravelling the complex interactions between the Earth’s subsystems, often termed as “spheres”: atmosphere, biosphere, hydrosphere (including oceans and cryosphere), pedosphere, or lithosphere, and increasingly the “anthroposphere”. The grand opportunity today is that many key processes in various subsystems of the Earth are constantly monitored. Networks of ecological, hydrometeorological, and atmospheric in situ measurements, for instance, provide continuous insights into the dynamics of the terrestrial water and carbon fluxes (Dorigo et al., 2011; Baldocchi, 2014; Wingate et al., 2015; Mahecha et al., 2017). Earth observations retrieved from satellite remote sensing enable a synoptic view of the planet and describe a wide range of phenomena in space and time (Pfeifer et al., 2012; Skidmore et al., 2015; Mathieu et al., 2017). The subsequent integration of in situ and space-derived data, e.g. via machine learning methods, leads to a range of unprecedented quasi-observational data streams (e.g. Tramontana et al., 2016; Balsamo et al., 2018; Bodesheim et al., 2018; Jung et al., 2019). Likewise, diagnostic models that encode basic process knowledge, but which are essentially driven by observations, produce highly relevant data products (see, e.g. Duveiller and Cescatti, 2016; Jiang and Ryu, 2016a; Martens et al., 2017; Ryu et al., 2018). Many of these derived data streams are essential for monitoring the climate system including land surface dynamics (see, for instance, the essential climate variables, ECVs; Hollmann et al., 2013; Bojinski et al., 2014), oceans at different depths (essential ocean variables, EOVs; Miloslavich et al., 2018), or the various aspects of biodiversity (essential biodiversity variables, EBVs; Pereira et al., 2013). Together, these essential variables describe the state of the planet at a given moment in time and are indispensable for evaluating Earth system models (Eyring et al., 2019).

With regard to the acquisition of sensor measurements and the derivation of downstream data products, Earth system sciences are well prepared. But can this multitude of data streams be used efficiently to diagnose the state of the Earth system? In principle, our answer would be affirmative, but in practical terms we perceive high barriers to interconnecting multiple data streams and further linking these to data analytic frameworks (as discussed for the EBVs by Hardisty et al., 2019). Examples of these issues are (i) insufficient data discoverability, (ii) access barriers, e.g. restrictive data use policies, (iii) lack of capacity building for interpretation, e.g. understanding the assumptions and suitable areas of application, (iv) quality and uncertainty information, (v) persistence of data sets and evolution of maintained data sets, (vi) reproducibility for independent researchers, (vii) inconsistencies in naming or unit conventions, and (viii) co-interpretability, e.g. either due to spatiotemporal alignment issues or physical inconsistencies, among others. Some of these issues are relevant to specific data streams and scientific communities only. In most cases, however, these issues reflect the neglect of the FAIR principles (to be “findable, accessible, interoperable, and re-usable”; Wilkinson et al., 2016). If the lack of FAIR principles and limited (co-)interpretability come together, they constitute a major obstacle in science and slow down the path to new discoveries. Or, to put it as a challenge, we need new solutions that minimize the obstacles that hinder scientists from capitalizing on the existing data streams and accelerate scientific progress. More specifically, we need interfaces that allow for interacting with a wide range of data streams and enable their joint analysis either locally or in the cloud.

As long as we do not overcome data interoperability limitations, Earth system sciences cannot fully exploit the promises of novel data-driven exploration and modelling approaches to answer key questions related to rapid changes in the Earth system (Karpatne et al., 2018; Bergen et al., 2019; Camps-Valls et al., 2019; Reichstein et al., 2019). A variety of approaches have been developed to interpret Earth observations and big data in the Earth system sciences in general (for an overview, see, e.g. Sudmanns et al., 2019) and gridded spatiotemporal data as a special case (Nativi et al., 2017; Lu et al., 2018). For the latter, data cubes have recently become popular, addressing an increasing demand for efficient access, analysis, and processing capabilities for high-resolution remote sensing products. The existing data cube initiatives and concepts (e.g. Baumann et al., 2016; Lewis et al., 2017; Nativi et al., 2017; Appel and Pebesma, 2019; Giuliani et al., 2019) vary in their motivations and functionalities. Most of the data cube initiatives are, however, motivated by the need for accessing singular (very-)high-resolution data cubes, e.g. from satellite remote sensing or climate reanalysis, and not by the need for global multivariate data exploitation.

This paper has two objectives: first, we aim to formalize the idea of an Earth system data cube (ESDC) that is tailored to explore a variety of Earth system data streams together and thus largely complements the existing approaches. The proposed mathematical formalism intends to illustrate
how one can efficiently operate such data cubes. Second, the paper aims at introducing the Earth System Data Lab (ESDL; https://earthsystemdatalab.net, last access: 21 February 2020). The ESDL is an integrated data and analytical hub that curates a multitude of data streams representing key processes of the different subsystems of the Earth in a common data model and coordinate reference system. This infrastructure enables researchers to apply their user-defined functions (UDFs) to these analysis-ready data (ARD). Together, these elements minimize the hurdle to co-explore a multitude of Earth system data streams. Most known initiatives intend to preserve the resolutions of the underlying data and facilitate their direct exploitation, like the Earth Server (Baumann et al., 2016) or the Google Earth Engine (Gorelick et al., 2017). The ESDL, instead, is built around singular data cubes on common spatiotemporal grids that include a high number of variables as a dimension in its own right. This design principle is thought to be advantageous compared to building data cubes from individual data streams without considering their interactions from the very beginning. Due to its multivariate structure and the easy-to-use interface, the ESDL is well suited for being part of data-driven challenges, as regularly organized by the machine learning community, for example.

The remainder of the paper is organized as follows: Sect. 2 introduces the concept based on a formal definition of Earth system data cubes and explains how user-defined functions can interact with them. In Sect. 3, we describe the implementation of the Earth System Data Lab in the programming language Julia and as a cloud-based data hub. Section 4 then illustrates three research use cases that highlight different aspects of the ESDL. We present an example from an univariate analysis, characterizing seasonal dynamics of some selected variables; an example from high-dimensional analysis; and an example for the representation of a model–data integration approach. In Sect. 5, we discuss the current advantages and limitations of our approach and put an emphasis on required future developments.

## 2 Concept

Our vision is that multiple spatiotemporal data streams shall be treated as a singular yet potentially very high-dimensional data stream. We call this singular data stream an Earth system data cube. For the sake of clarity, we introduce a mathematical representation of the Earth system data cube and define operations on it. Further details on an efficient implementation are provided in Sects. 3.2 and 3.3.

Suppose we observe \( p \) variables \( Y^1, \ldots, Y^p \), each under a (possibly different) range of conditions. A first step towards data integration is to (re)sample all data streams onto a common domain \( J \) (e.g. a spatiotemporal grid) to obtain the indexed set \( \{Y^1_j, \ldots, Y^p_j\}_{j \in J} \) of multivariate observations. Observations obtained from different variables are then identified as different coordinates in the multivariate array \( Y \). However, when calculating simple variable summaries or performing spatiotemporal aggregations of the data, such a representation can be computationally obstructive. We therefore propose to consider the “variable indicator” \( k \in \{1, \ldots, p\} \) as simply another dimension of the index set and view the data as the collection \( \{X^k_j\}_{j \in J} \) of univariate observations, where \( I = J \times \{1, \ldots, p\} \) and where \( X^k_{i,j} := Y^k_j \). With this idea in mind, we now formally define the Earth system data cube (“data cube” in short).

A data cube \( C \) consists of a triplet \( (L, G, X) \) of components to be described below.

- \( L \) is a set of labels, called dimensions, describing the axes of the data cube. For example, \( L = \{\text{lat, long, time, var}\} \) describes a data cube containing spatiotemporal observations from a range of different variables. The number of dimensions \( |L| \) is referred to as the order of cube \( C \); in the above example, \( |L| = 4 \).

- \( G \) is a collection \( \{\text{grid}(\ell)\}_{\ell \in L} \) of grids along the axes in \( L \). For every \( \ell \in L \), the set \( \text{grid}(\ell) \) is a discrete subset of the domain of the axis \( \ell \), specifying the resolution at which data are available along this axis. Every set \( \text{grid}(\ell) \) is required to contain at least two elements. Dimensions containing only one grid point are dropped. The collection \( G \) defines the hyper-rectangular index set \( I(G) := \times_{\ell \in L} \text{grid}(\ell) \), motivating the name “cubical” for \( C \). For example,

\[
I(G) = \times_{\ell \in L} \text{grid}(\ell)
= \text{grid(lat)} \times \text{grid(long)} \times \text{grid(time)} \times \text{grid(var)}
= \{(-89.75, \ldots, 89.75) \times (-179.75, \ldots, 179.75) \times [1 \text{ Jan 2010}, \ldots, 31 \text{ Dec 2010}] \times \{\text{GPP, SWC, } R_g\} \}
= \{(-89.75, -179.75, 1 \text{ Jan 2010, GPP}),
\ldots, (89.75, 179.75, 31 \text{ Dec 2010, } R_g)\}\).
\]

Since \( G \) and \( I(G) \) are in one-to-one correspondence, we will use the two interchangeably.

- \( X \) is a collection of data \( \{X^k_i\}_{i \in I(G)} \subseteq \mathbb{R}_{\text{NA}} := \mathbb{R} \cup \{\text{NA}\} \) observed at the grid points in \( I(G) \). Here, “NA” refers to “not available”.

In this view, the data can be treated as a collection \( \{X^k_i\}_{i \in I(G)} \) of univariate observations, even if they encode different variables. In the above example, the variable axis is a nominal grid with the entries GPP (gross primary production), SWC (soil water content), and \( R_g \) (global radiation). The set of all data cubes with dimensions \( L \) will be denoted by \( C(L) \). Data cubes that contain one variable only can be considered as special case; other common choices of \( L \) are described in Table 1. The list of example axes labels used in

\(^1\)The symbol \( \times \) indicates a Cartesian product.
the table is, of course, not exhaustive. Other relevant dimensions could be, for example, model versions, model parameters, quality flags, or uncertainty estimates. Note that, by definition, a data cube only depends on its dimensions through the set of axes $L$ and is therefore indifferent to any order of these. In the remainder of this article, the notion of data cubes refers to this concept. Note that dropping dimensions that only contain one grid point is not the only possible way of working with data cubes. Another equally valid idea is to maintain grids of length 1 and integrate them to the workflow.

2.1 Operations on an Earth system data cube

To exploit an Earth system data cube efficiently, scientific workflows need to be translated into operations executable on data cubes as described above. More specifically, the output of each operation on a data cube should yield another data cube. The entire workflow of a project, possibly a succession of analyses performed by different collaborators, can then be expressed as a composition of several UDFs performed on a single (input) data cube. Besides unifying all statistical data analyses into a common concept, the idea of expressing workflows as functional operations on data cubes comes with another important advantage: as soon as a workflow is implemented as a suitable set of UDFs, it can be reused on any other sufficiently similar data cube to produce the same kind of output.

In its most general form, a user-defined function $f : C \to C$ operates by (i) extracting relevant information from $C$, (ii) performing calculations on the extracted information, and (iii) storing these calculations into a new data cube $f(C)$. In order to perform step (i), $f$ expects a minimal set of dimensions $E$ of the input cube. The returned set of axes for an input cube with dimensions $E$ will be denoted by $R$. That is, $f$ is a mapping such that

$$f : C(E) \to C(R).$$

Alongside the function $f$, one has to define the sets $E$ and $R$, which we will refer to as minimal input and minimal output dimensions, respectively.

A major advantage of thinking in data cube workflows is that low-dimensional functions can be applied to higher-dimensional cubes by simple functional extensions: a function can be acting along a particular set of dimensions while looping across all unspecified dimensions. For example, the function that computes the temporal mean of a univariate time series should allow for an input data cube, which, in addition to a temporal grid, contains spatial information. The output of such an operation should then be a cube of spatially gridded temporal means. Similarly, the function should be applicable to cubes containing multivariate observations. Here, we expect the output to contain one temporal mean per supplied variable.

In general, a function $f$ defined on $C(E)$ should naturally extend to a function from $C(E \cup A)$ to $C(R \cup A)$ with $A \cap R = \emptyset$ by executing the described “apply” operation. The code package accompanying this paper (described in Sect. 3) automatically equips every UDF with such a functionality. A schematic description of this approach is illustrated in Fig. 1.

The approach outlined above is very convenient to describe workflows, i.e. recursive chains of UDFs. Let $f_1, \ldots, f_n$ be a sequence of UDFs with corresponding minimal input/output dimensions ($E_1, R_1), \ldots, (E_n, R_n)$. If an output dimension $R_i$ is a subset of subsequent input $E_{i+1}$, we can chain these functions. A recursive workflow emerges when $R_i \subseteq E_{i+1}$ for all $i$, by iteratively chaining $f_1, \ldots, f_n$ upon one another. The input/output dimensions of the resulting cube are $(E_1, R_n)$.

Overall, the definition of an Earth system data cube and associated operations on it do not only guide the implementation strategy but also help us summarize potentially complicated analytic procedures in a common language. For the sake of readability, in the following, we will not distinguish between a function $f$ (defined only for minimal input) and its extension $\tilde{f}$ (equipped with the apply functionality; see Fig. 1). The former will be referred to as an “atomic” function. We typically indicate the minimal input/output dimensions $(E, R)$ of a function $f$ by writing $f^E_E$. Since the pair $(E, R)$ does not determine the mapping $f$, this notation should not be understood as the parameterization of a function class but rather provide an easy way to perform input control and to anticipate the output dimensions of a cube returned by $f$. For instance, following the discussion above, a function denoted by $f^E_E$ can be applied to any cube with dimension $E \cup A$, satisfying that $A \cap R = \emptyset$, and returns a cube with dimensions $R \cup A$. To avoid ambiguities, additional notation is needed when distinguishing between two functions with the same pair of minimal input/output dimensions.

2.2 Examples

In the following, we present some special operations that are routinely needed in explorations of Earth system data cubes: “Reducing” describes a function that calculates some scalar measure (e.g. the sample mean). Consider, for instance, the need to estimate the mean of a univariate data cube, of course weighted by the area of the spatial grid cells. An operation of this kind expects a cube with dimensions $E = \{\text{lat}, \text{long}, \text{time}\}$ and returns a cube with dimensions $R = \{\}$ and is therefore a mapping:

$$f^{(1)}_{\text{lat,long,time}} : C(\{\text{lat, long, time}\}) \to C(\{\}).$$

This mapping can now be applied to any data cube of potentially higher (but not lower) dimensionality. For instance, $f$ is automatically extended to a multivariate spatiotemporal data cube (Table 1) with the mapping

$$f^{(1)}_{\text{lat,long,tim, var}} : C(\{\text{lat, long, time, var}\}) \to C(\{\text{var}\}),$$

which computes one spatiotemporal mean for each variable.
Table 1. Typical sets of data cubes $C(L)$ of varying orders $|L|$ with characteristic dimensions $L$.

| Order $|L|$ | Set of data cubes $C(L)$ | Description of $C(L)$ |
|--------|--------------------------|------------------------|
| 0      | $C(\emptyset)$          | Scalar value where no dimension is defined |
| 1      | $C(\{\text{lat}\})$     | Univariate latitudinal profile |
| 1      | $C(\{\text{long}\})$    | Univariate longitudinal profile |
| 1      | $C(\{\text{time}\})$    | Univariate time series |
| 1      | $C(\{\text{var}\})$     | Single multivariate observation |
| 2      | $C(\{\text{lat, long}\})$ | Univariate static geographical map |
| 2      | $C(\{\text{lat, time}\})$ | Univariate Hovmöller diagram: zonal pattern over time |
| 2      | $C(\{\text{lat, var}\})$ | Multivariate latitudinal profile |
| 2      | $C(\{\text{long, time}\})$ | Univariate Hovmöller diagram: meridional pattern over time |
| 2      | $C(\{\text{long, var}\})$ | Multivariate longitudinal profile |
| 2      | $C(\{\text{time, var}\})$ | Multivariate time series |
| 2      | $C(\{\text{time, freq}\})$ | Univariate time–frequency plane |
| 3      | $C(\{\text{lat, long, time}\})$ | Univariate data cube |
| 3      | $C(\{\text{lat, long, var}\})$ | Multivariate map, e.g. a global map of different soil properties |
| 3      | $C(\{\text{lat, time, var}\})$ | Multivariate latitudinal Hovmöller diagram |
| 3      | $C(\{\text{long, time, var}\})$ | Multivariate longitudinal Hovmöller diagram |
| 3      | $C(\{\text{time, freq, var}\})$ | Multivariate spectrally decomposed time series |
| 4      | $C(\{\text{lat, long, time, var}\})$ | Multivariate spatiotemporal cube |
| 4      | $C(\{\text{lat, long, time, freq}\})$ | Univariate spectrally decomposed data cube |
| 5      | $C(\{\text{lat, long, time, var, ens}\})$ | Multivariate ensemble of model simulations |

Figure 1. Schematic illustration of the “apply” functionality: a function $f : C(E) \to C(R)$ is extended to the set of cubes with dimensions $E \cup A$, where $A$ is an arbitrary set of dimensions with $A \cap R = \emptyset$. Given a cube $C \in C(E \cup A)$, the extension $\overline{f}(C)$ is constructed by iterating over all grid points $i$ along the dimensions in $A$ to obtain the collection $\{C_i\} \subseteq C(E)$ of sliced cubes, applying $f$ to every cube $C_i$ separately, and binding the collection $\{f(C_i)\}$ into the output cube $\overline{f}(C) \in C(R \cup A)$. Here, the index $i$ runs through all elements in $\times_{\text{grid}(a)}$. 
“Cropping” is subsetting a data cube while maintaining the order of a cube. A cropping operation typically reduces certain axes of a data cube to only contain specified grid points (and therefore requires the input cube to contain these grid points). For instance, a function that extracts a certain “cropped” fraction \( T_0 \) along the temporal axis expects an input cube containing a time axis with a grid at least as highly resolved as \( T_0 \). This function preserves the dimensionality of the cube but reduces the grid along the time axis; i.e.

\[
 f_{\{\text{time}\}}^{\{\text{time}\}} : \mathcal{C}((\text{time})|\text{grid}(\text{time}) \geq T_0) \rightarrow \mathcal{C}((\text{time})|\text{grid}(\text{time}) = T_0),
\]

where we have used \( \mathcal{C}(L|P) \) to denote the set of cubes with dimensions \( L \) satisfying the condition \( P \). Thanks to the apply functionality, this atomic function can be used on any cube of higher order. For example, it is readily extended to a mapping:

\[
 f_{\{\text{time}\}}^{\{\text{time}\}} : \mathcal{C}([\text{lat}, \text{long}, \text{time}]|\text{grid}(\text{time}) \geq T_0) \rightarrow \mathcal{C}([\text{lat}, \text{long}, \text{time}]|\text{grid}(\text{time}) = T_0),
\]

which crops the time axis of cubes with dimensions \([\text{lat}, \text{long}, \text{time}]\). Analogously, all dimensions can be subsetted as long as the length of the dimension is larger than 1. The latter would be called slicing.

“Slicing” refers to a subsetting operation in which a dimension of the cube is degenerated, and the order of the cube is reduced and can be interpreted as a special form of cropping. For instance, if we only select a singular time instance \( t_0 \), the time dimension effectively vanishes as we do not longer need a vector-spaced dimension to represent its values. When applied to a spatiotemporal data cube, this amounts to a mapping:

\[
 f_{\{\text{time}\}}^{\{\text{time}\}} : \mathcal{C}([\text{lat}, \text{long}, \text{time}]|\text{grid}(\text{time}) \geq t_0) \rightarrow \mathcal{C}([\text{lat}, \text{long}]),
\]

“Expansions” are operations where the order of the output cube is higher than the order of the corresponding input cube. Thanks to the apply functionality, this atomic function can be used on any cube of higher order. For example, it is readily extended to a mapping:

\[
 f_{\{\text{time}\}}^{\{\text{time}\}} : \mathcal{C}((\text{time})|\text{grid}(\text{time}) \geq t_0) \rightarrow \mathcal{C}((\text{time}), \text{freq}).
\]

“Multiple cube handling” is often needed, for instance, when fitting a regression model where response and predictions are stored in different cubes. Also, we may be interested in outputting the fitted values and the residuals in two separate cubes. This amounts to an atomic operation:

\[
 f_{\{\text{para}, \text{time}\}}^{\{\text{para}, \text{time}\}} : \mathcal{C}((\text{time}, \text{var}) \times \mathcal{C}((\text{time}))) \rightarrow \mathcal{C}((\text{para}) \times \mathcal{C}((\text{time})),
\]

which expects a multivariate data cube for the predictors \( C_1 \in \mathcal{C}((\text{time}, \text{var})) \) and a univariate cube for the targets \( C_2 \in \mathcal{C}((\text{time})) \). The output consists of a vector of fitted parameters \( C_1 \in \mathcal{C}((\text{para})) \) and a residual time series \( C_2 \in \mathcal{C}((\text{time})) \) to compute the model performance. This concept also allows the integration of more than two input and/or output cubes.

3 Data streams and implementation

The concept as described in Sect. 2 is generic, i.e. independent of the implemented Earth system data cube and of the technical solution of the implementation. Figure 2 shows how the concept outlined above is realized from a practical point of view. The flowchart shows that the starting point is the collection of relevant data streams which then need to be preprocessed in order to be interpretable as a single data cube. The ESDC itself may be stored locally or in the cloud and can be accessed from various users simultaneously based on different application programming interfaces (APIs). In the following, we firstly present the data used in our implementation of the ESDL which is available online, and secondly describe the implementation strategy for the API we developed in this project.

3.1 Data streams in the ESDL

The data streams included so far were chosen to enable research on the following topics (a complete list is provided in Appendix A):

- Ecosystem states at the global scale in terms of relevant biophysical variables. Examples are leaf area index (LAI), the fraction of photosynthetically active radiation (fAPAR), and albedo (Disney et al., 2016; Pinty et al., 2006; Blessing and Löw, 2017).

- Biosphere–atmosphere interactions as encoded in land fluxes of \( \text{CO}_2 \), i.e. GPP, terrestrial ecosystem respiration \( (R_{\text{eco}}) \), and the net ecosystem exchange \( \text{(NEE)} \) as well as the latent heat \( (LE) \) and sensible heat \( (H) \) energy fluxes. Here, we rely mostly on the FLUXCOM data suite (Tramontana et al., 2016; Jung et al., 2019).

- Terrestrial hydrology requires a wide range of variables. We mainly ingest data from the Global Land Evaporation Amsterdam Model (GLEAM; Martens et al., 2017; Miralles et al., 2011) which provide a series of relevant surface hydrological properties such as surface (SM) and root-zone soil moisture \( (\text{SM}_{\text{root}}) \) but also potential evaporation \( (E_p) \) and evaporative stress \( (S) \) conditions, among others. Ingesting entire products such as GLEAM ensures internal consistency.

- State of the atmosphere is described using data generated by the Climate Change Initiative (CCI) by the European Space Agency (ESA) in terms of aerosol optical depth at different wavelengths \( (\text{AOD}_{550}, \text{AOD}_{555}, \text{AOD}_{659}, \text{and AOD}_{1610}) \); Holzer-Popp et al., 2013), total ozone column \( (\text{Van Roozendael et al., 2012; Lerot et al., 2014}) \), as well as surface ozone (which is more relevant to plants), and total column water vapour \( (\text{TCWV}; \text{Schröder et al., 2012; Schneider et al., 2013}) \).
Figure 2. Workflow putting the ESDL concept into practice: selected data sets are preprocessed to common grids and saved in cloud-ready data formats (Zarr). Based on these cubed data sets, a global Earth system data cube can be produced that is either stored locally or in the cloud. Via appropriate application programming interfaces (APIs), users can efficiently access the ESDC in their native language. Users can fully focus on designing user-defined functions and workflows.

- Meteorological conditions are described via the reanalysis data, i.e. the ERA5 product. Additionally, precipitation is ingested from the Global Precipitation Climatology Project (GPCP; Adler et al., 2003; Huffman et al., 2009).

Together, these data streams form data cubes of intermediate spatial and temporal resolutions (0.25, 0.083°; both 8 d), visualized in Fig. 3. The variables described here are described in more detail in a list provided in Appendix A, which may, however, already be incomplete at the time of publication, as the ESDL is a living data suite, constantly expanding according to users’ requests. For the latest overview, we refer the reader to the website (https://www.earthsystemdatalab.net/, last access: 21 February 2020). Note that we have not considered the integration of uncertainty as another dimension in the current implementation. The rationale is that each of the data products comes with a specific uncertainty flag or estimate that cannot be merged in an own dimension. This is an open aspect that needs to be addressed in future developments.

To show the portability of the approach, we have developed a regional data cube for Colombia. This work supports the Colombian Biodiversity Observational Network activities within the Group on Earth Observations Biodiversity Observation Network (GEO BON). This regional data cube has a 1 km (0.083°) resolution and focuses on remote-sensing-derived data products (i.e. LAI, FAPAR, the normalized difference vegetation index (NDVI), the enhanced vegetation index (EVI), LST, and burnt area). In addition to the global ESDL, monthly mean products such as cloud cover (Wilson and Jetz, 2016) have been ingested given their recurrent applicability in biodiversity studies at regional scales. Data layers from governmental organizations providing detailed information about ecosystems are also available that allow a national characterization and deeper understanding of ecosystem changes by natural or human drivers. These are maps of biotic units (Londoño et al., 2017), wetlands (Florez et al., 2016), and agriculture frontier maps (MADR-UPRA, 2017). Additionally, GPP, evapotranspiration, shortwave radiation, PAR, and diffuse PAR from the Breathing Earth System Simulator (BESS; Ryu et al., 2011, 2018; Jiang and Ryu, 2016b) and albedo from QA4ECV (http://www.qa4ecv.eu/, last access: 21 February 2020) are available, among others. This regional Earth system data cube should serve as a platform for analysis in a region with variability of landscape, high biodiversity and ecosystem transitions gradients, and facing rapid land use change (Sierra et al., 2017).

3.2 Implementation

To put the concept of an Earth system data cube as outlined in Sect. 2 into practice, we need suitable access APIs (see Fig. 2). A co-author of this paper (FG) developed one API in the relatively young scientific programming language Julia (https://julialang.org/, last access: 21 February 2020; Bezanson et al., 2017) which is provided via the ESDL.jl package. Additionally, all functionalities are also available in Python based on existing libraries and documented online. In both cases, the goal was that the user does not have to explicitly deal with the complexities of sequential data input/output handling and can concentrate on implementing the atomic functions and workflows, while the system takes care of necessary out of core and out-of-memory computations.
Figure 3. Visualization of the implemented Earth system data cube (an animation is provided online at https://youtu.be/9L4-fq48Ev0, last access: 21 February 2020). The figure shows from the top left to bottom right the variables sensible heat ($H$), latent heat ($LE$), gross primary production (GPP), surface moisture (SM), land surface temperature (LST), air temperature ($T_{air}$), cloudiness ($C$), precipitation ($P$), and water vapour ($V$). References to the individual data sources are given in Appendix A. Here, the resolution in space is $0.25^\circ$ and 8 d in time, and we are inspecting the time from May 2008 to May 2010; the spatial range is from 15$^\circ$ S to 60$^\circ$ N, and 10$^\circ$ E to 65$^\circ$ W.

The following is a sketched description of the principles of the Julia-based ESDL.jl implementation. We chose Julia to translate the concepts outlined into efficient computer code because it has clear advantages for data cube applications besides its general elegance in scientific computing in terms of speed, dynamic programming, multiple dispatch, and syntax (Perkel, 2019). Specifically, Julia allows for generic processing of high-dimensional data without large code repetitions. At the core of the Julia ESDL.jl toolbox are the mapslices and mapCube functions, which execute user-defined functions on the data cube as follows:

- Given some large data cube $C = (L, G, X)$, the ESDL function `subsetcube(C)` will retrieve a handle to $C$ that fully describes $L$ and $G$.

- Knowledge of the desired $L$ and $G$ allows us to develop a suitable user-defined function $f^E_R$.

- Depending on the exact needs, mapslices and mapCube will then be used to apply the $f^E_R$ on a cube as illustrated in Fig. 1. mapCube is a strict implementation of the cube mapping concept described here, where it is mandatory to explicitly describe $E$ and $R$ such that the atomic function is fully operational.

The case studies developed in Sect. 4 are accompanied by code that illustrates this approach in practice. Of course there are also alternatives to Julia. Lu et al. (2018) recently reviewed different ways of applying func-
tions on array data sets in R, Octave, Python, Rasdaman, and SciDB. One requirement of such a mapping function is that it should be scalable, which means that it should process data larger than the computer memory and, if needed, in parallel. While existing solutions are sufficient for certain applications, most are not consistent with the cube mapping concept as described in Sect. 2. For instance, the required handling of complex workflows of multiple cubes (Eq. 8) is typically not possible in the existing solutions that have been reviewed. In some cases, issues in the computational efficiency of the underlying programming languages render certain solutions not suitable. This is particularly the case when user-defined functions become complex. Likewise, certain properties such as the desired indifference to the ordering in axes dimensions are often not foreseen. One suitable alternative to Julia is available in Python. The xarray (http://xarray.pydata.org, last access: 21 February 2020) and dask packages have been successfully utilized in the Open Data Cube, Pangeo, and xcube initiatives. Extensive descriptions on how to work in the ESDL with both Python and Julia can be accessed from the following website: https://www.earthsystemdatalab.net/ (last access: 21 February 2020). The open-source implementation of the ESDL also implies that one can easily extend the stored data sets. The online documentation shows in detail how additional data can be added to the ESDL. In particular, if the data share common axes and are stored in a compatible format (as described below in Sect. 3.3), this does not require major efforts.

3.3 Storage and processing of the data cube

The ESDL has been built as a generic tool. It is prepared to handle very large volumes of data. Storage techniques for large raster geodata are generally split into two categories: database-like solutions like Rasdaman (Baumann et al., 1998) or SciDB (Stonebraker et al., 2013) access data directly through file formats that follow metadata conventions like HDF5 (https://www.hdfgroup.org/, last access: 21 February 2020) or NetCDF (https://www.unidata.ucar.edu/software/netcdf/, last access: 21 February 2020). Database solutions shine in settings where multiple users repeatedly request (typically small) subsets of data cube, which might not be rectangular, because the database can accelerate access by adjusting to common access patterns. However, for batch processing large portions of a data cube, every data entry is ideally accessed only once during the whole computation. Hence, when large fractions of some data cube have to be accessed, users will usually avoid the overhead of building and maintaining a database and rather aim for accessing the data directly from their files. This experience is often perceived as more “natural” for Earth system scientists who are used to “touching” their data, knowing where files are located, and so forth. Databases instead offer, by construction, an entry point to an otherwise unknown data set.

One disadvantage of the traditional file formats used for storing gridded data is that their data chunks are contained in single files that may become impossible to handle efficiently. This is not problematic when the data are stored on a regular file system where the file format library can read only parts of the file. In cloud-based storage systems, it is not common to have an API for accessing only parts of an object, so these file formats are not well suited for being stored in the cloud. Recently, novel solutions for this issue were proposed, including modifications to existing storage formats, e.g. HDF5 cloud, or cloud-optimized GeoTiff, among others, as well as completely new storage formats, in particular Zarr (https://zarr.readthedocs.io/, last access: 21 February 2020) and TileDB (https://tiledb.io/, last access: 21 February 2020). While working with these formats is very similar to traditional solutions (like HDF5 and NetCDF), these new formats are optimized for cloud storage as well as for parallel read and write operations. Here, we chose to use the new Zarr format. The reason is that it enables us to share the data cube through an object storage service, where the data are public and can be analysed directly. Python packages for accessing and analysing large N-dimensional data sets like xarray and dask, which make a wide range of existing tools readily usable on the cube, and a Julia approach to read Zarr data have been implemented as well.

At present, the ESDL provides the same data cube in different spatial resolutions and different chunkings to speed up data access for different applications. In chunked data formats, a large data set is split into smaller chunks, which can be seen as separate entities where each chunk is represented by an object in an object store. There are several ways to chunk a data cube. Consider the case of a multivariate spatiotemporal cube \( C(l_{\text{at}}, l_{\text{ong}}, t_{\text{ime}}, v_{\text{ar}}) \). One common strategy would be to treat every spatial map of each variable and time point as one chunk, which would result in a chunk size of \([\text{grid}(l_{\text{at}}) \times \text{grid}(l_{\text{ong}})] \times 1 \times 1\). However, because an object can only be accessed as a whole, the time for reading a slice of a univariate data cube does not directly scale with the number of data points accessed but rather with the number of accessed chunks. Reading out a univariate time series of length 100 from this cube would require accessing 100 chunks. If one stored the same data cube with complete time series contained in one chunk, read operations could perform much faster. Table 2 shows an overview of the implemented chunkings for different cubes in the current ESDL environment.

4 Experimental case studies

The overarching motivation for building an Earth system data cube is to support the multifaceted needs of Earth system sciences. Here, we briefly describe three case studies of varying complexity (estimating seasonal means per latitude, dimensionality reduction, and model–data integration) to illu-
Table 2. Resolutions and chunkings of the currently implemented global Earth system data cube per variable. Here, the cubes with chunk size 1 in the time coordinate are optimized for accessing global maps at a time, while the other cubes are more suited for processing time series or regional subsets of the data cube. The cubes are currently hosted on the Object Storage Service by the Open Telecom Cloud under https://obs.eu-de.otc.t-systems.com/obs-esdc-v2.0.0/ (last access: 21 February 2020) (state: September 2019).

<table>
<thead>
<tr>
<th>Resolution</th>
<th>Chunk size along axis</th>
<th>Grid (time)</th>
<th>Grid (lat)</th>
<th>Grid (long)</th>
</tr>
</thead>
<tbody>
<tr>
<td>0.083°</td>
<td>1</td>
<td>170</td>
<td>90</td>
<td>90</td>
</tr>
<tr>
<td>0.083°</td>
<td>2</td>
<td>2160</td>
<td>90</td>
<td>90</td>
</tr>
<tr>
<td>0.25°</td>
<td>1</td>
<td>170</td>
<td>90</td>
<td>1440</td>
</tr>
</tbody>
</table>

4.1 Inspecting summary statistics of biosphere–atmosphere interactions

Data exploration in the Earth system sciences typically starts with inspecting summary statistics. Global mean patterns across variables can give an impression on the long-term system behaviour across space. In this first use case, we aim to describe mean seasonal dynamics of multiple variables across latitudes.

Consider an input data cube of the form \( C(\{\text{lat}, \text{long}, \text{time}, \text{var}\}) \). The first step consists in estimating the median seasonal cycles per grid cell. This operation creates a new dimension encoding the “day of year” (doy) as described in the atomic function of Eq. (9):

\[
\mathcal{f}_{\text{doy}}^{\text{time}} : C(\{\text{lat}, \text{long}, \text{time}, \text{var}\}) \rightarrow C(\{\text{lat}, \text{long}, \text{doy}, \text{var}\}). \tag{9}
\]

In a second step, we apply an averaging function that summarizes the dynamics observed at all longitudes:

\[
\mathcal{f}_{\text{long}}^{(1)} : C(\{\text{lat}, \text{long}, \text{doy}, \text{var}\}) \rightarrow C(\{\text{lat}, \text{doy}, \text{var}\}). \tag{10}
\]

The result is a cube of the form \( C(\{\text{lat}, \text{doy}, \text{var}\}) \) describing the seasonal pattern of each variable per latitude. Figure 4 visualizes this analysis for data on GPP, air temperature (\( T_{\text{air}} \)), and surface moisture (SM; all references for data streams used are provided in Appendix A). The first row visualizes GPP; on the left side, we see the Northern Hemisphere, where darker colours describe higher latitudes and the background is the actual value of the variable. Together, the left and right plots describe the global dynamics of phenology, often referred to as the “green wave” (Schwartz, 1998). We clearly see the almost nonexistent GPP in high-latitude winters and also find the imprint of constantly low to intermediate productivity values at latitudes that are characterized by dry ecosystems. Pronounced differences between Northern and Southern Hemisphere reflect the very different distribution of productive land surface.

For temperature, the observed seasonal dynamics are less complex. We essentially find the constantly high temperature conditions near the Equator and visualize the pronounced seasonality at high latitudes. However, Fig. 4 also shows that temperature peaks lag behind the June/December solstices in the Northern Hemisphere, while in the Southern Hemisphere, the asymmetry of the seasonal cycle in temperature is less pronounced. While the seasonal temperature gradient is a continuum, surface moisture shows a much more complex pattern across latitudes, as reflected in summer/winter depressions in certain midlatitudes. For instance, a clear drop at, e.g., latitudes of approximately 60°N and even stronger depressions in latitudinal bands dominated by dry ecosystems.

This example analysis is intended to illustrate how the sequential application of two basic functions on this Earth system data cube can unravel global dynamics across multiple variables. We suspect that applications of this kind can lead to new insights into apparently known phenomena, as they allow to investigate a large number of data streams simultaneously and with consistent methodology.

4.2 Intrinsic dimensions of ecosystem dynamics

The main added value of the ESDL approach is its capacity to jointly analyse large numbers of data streams in integrated workflows. A long-standing question arising when a system is observed based on multiple variables is whether these are all necessary to represent the underlying dynamics. The question is whether the data observed in \( Y \in \mathbb{R}^M \) could be described with a vector space of much smaller dimensionality \( Z \in \mathbb{R}^m \) (where \( m \ll M \)), without loss of information, and what value this “intrinsic dimensionality” \( m \) would have (Lee and Verleysen, 2007; Camastra and Staiano, 2016). Note that in this context the term “dimension” has a very different connotation compared to the “cube dimensions” introduced above.

When thinking about an Earth system data cube, the question about its intrinsic dimensionality could be investigated along the different axes. In this study, we ask if the multitude of data streams, grid(\text{var}), contained in our Earth system data cube is needed to grasp the complexity of the terrestrial surface dynamics. If the compiled data streams were highly redundant, it could be sufficient to concentrate on only a few orthogonal variables and design the development of the study accordingly. Starting from a cube \( C(\{\text{lat}, \text{long}, \text{time}, \text{var}\}) \),
we ask at each geographical coordinate if the local vector space spanned by the variables can be compressed such that $m_{\text{var}} \ll |\text{grid}(\text{var})|$.

Estimating the intrinsic dimension of high-dimensional data sets has been a matter of research for multiple decades, and we refer the reader to the existing reviews on the subject (e.g. Camastra and Staiano, 2016; Karbauskaite and Dzemyda, 2016). An intuitive approach is to measure the compressibility of a data set via dimensionality reduction techniques (see, e.g. van der Maaten et al., 2009; Kraemer et al., 2018). In the simplest case, one can apply a principal component analysis (PCA, using different time points as different observations) and estimate the number of components that together explain a predefined threshold of the data variance. In our application, we followed this approach and chose a threshold value of 95% of variance. The atomic function needed for this study is described in Eq. (11):

$$f^{(1)}_{\text{time, var}}: C((\text{lat}, \text{long}, \text{time}, \text{var})) \rightarrow C((\text{lat}, \text{long})).$$  (11)
Figure 5. Intrinsic dimension of 18 land ecosystem variables. The intrinsic dimension is estimated by counting how many principal components would be needed to explain at least 95% of the variance in the Earth system data cube. The results for the original data are shown in panel (a). The analysis is then repeated based on subsignals of each variable, representing different timescales. In panel (b), we show the intrinsic dimension of long-term modes of variability, in (c) for modes representing seasonal components, and (d) for modes of short-term variability. Light grey areas indicate zones where at least one data stream was incomplete and no intrinsic dimension could be estimated based on the same set of variables.

The output is a map of spatially varying estimates of intrinsic dimensions $m_{\text{var}}$. We performed this study considering the following 18 variables relevant to describing land surface dynamics: GPP, $R_{\text{eco}}$, NEE, $LE$, $H$, LAI, fAPAR, black- and white-sky albedo (each from two different sources), SM$_{\text{root}}$, $S$, transpiration, bare soil evaporation, evaporation, net radiation, and LST.

Figure 5 shows the results of this analysis for the original data, where the visualized range of intrinsic dimensions ranges from 2 to 13 (the analysis very rarely returns values of 1). At first glance, we find that ecosystems near the Equator are of higher intrinsic dimension (up to values of 12) compared to the rest of the land surface. In regions where we expect pronounced seasonal patterns, the intrinsic dimensionality is apparently low. We can describe these patterns by 4–7 dimensions. One explanation is that in cases where the seasonal cycle controls ecosystem dynamics, much of the surface variables tend to covary. This alignment implies that one can represent the dominant source of variance with few components of variability. In regions where the seasonal cycle plays only a marginal role, other sources of variability dominate that are, however, largely uncorrelated.

To verify that seasonality is the main source of variability in our analysis, we extend the workflow by decomposing each time series (by variable and spatial location) into a series of subsignals via a discrete fast Fourier transform (FFT). We then binned the subsignals into short-term, seasonal, and long-term modes of variability (as in Mahecha et al., 2010a; Linscheid et al., 2020), which leads to an extended data cube as we have shown in Eq. (12).

$$\mathcal{C}(\{\text{lat, long, time, var}\}) \rightarrow \mathcal{C}(\{\text{lat, long, time, var, freq}\})$$ (12)

The resulting cube is then further processed in Eq. (13) (which is the analogue to Eq. 11) to extract the intrinsic dimension per timescale:

$$\mathcal{C}(\{\text{lat, long, time, var, freq}\}) \rightarrow \mathcal{C}(\{\text{lat, long, freq}\})$$ (13)

The timescale-specific intrinsic dimension estimates only partly confirm the initial conjecture (Fig. 5). Short-term modes of variability always show relatively high intrinsic dimensions; i.e. the high-frequency components in the variables are rather uncorrelated. This finding can either be a
hint that we are seeing a set of independent processes or simply mean noise contamination. Seasonal modes, indeed, are of low intrinsic dimensionality, but considering that these modes are driven essentially by solar forcing only, they are surprisingly high dimensional. Additionally, we find a clear gradient from the inner tropics to arid and northernmost ecosystems. Warm and wet ecosystems seem to be characterized by a complex interplay of variables even when analysing their seasonal components only (see also Linscheid et al., 2020). One reason could be that seasonality in these regions is only marginally relevant to the total signal, or that tropical seasonality is inherently complicated. In the northern regions of South America, we find that arid regions seem to have low intrinsic seasonal dimensionality compared to more moist regions.

Long-term modes of land surface variability show a rather complex spatial pattern in terms of intrinsic dimensions: overall, we find values between 6 and 7 (see also the summary in Fig. 6). The values tend to be higher in high-altitude and tropical regions, whereas arid regions show low-complexity patterns. Long-term modes of variability in land surface variables are probably more complex than one would suspect a priori and should be analysed deeper in the near future.

The analysis shows how a large number of variables can be seamlessly integrated into a rather complex workflow. However, the results should be interpreted with caution: one criticism of the PCA approach is its tendency to overestimate the correct intrinsic dimensions in the presence of non-linear dependencies between variables. A second limitation is that the maximum intrinsic dimensions depend on the number of Fourier coefficients used to construct the signals, leading to different theoretical maximum intrinsic dimensions per timescale.

The question of the underlying dimensionality could also be investigated in a different way. While this study investigates the intrinsic dimensionality locally, i.e. along the dimensions of latitude and longitude, another recent study based on the ESDL by Kraemer et al. (2019) used a global PCA. Each observation is a point with coordinates “lat”, “long” and “time”, and the aim is to compress the “var” dimension. The form of the analysis is the following:

\[
\mathcal{C}([\text{lat, long, time, var}]) \rightarrow \mathcal{C}([\text{lat, long, time, princomp}]),
\]

and was applied to a subset of ESDL variables that describe dynamics in terrestrial ecosystems. This study corroborates the idea that land surface dynamics can be well represented in a surprisingly low-dimensional space. The analysis presented by Kraemer et al. (2019) suggests globally a much lower intrinsic dimensionality of 3 compared to what we find here based on a grid-cell-level analysis. This number corresponds to areas that are marked by a strong seasonality in our case. This is plausible, because the areas that show high intrinsic dimensionality in Fig. 5 are those where seasonal variability is low compared to the high-frequency variability (Linscheid et al., 2020). Local effects of this kind vanish when all spatial points are jointly analysed.

### 4.3 Model parameter estimation in the ESDL

Another key element in supporting Earth system sciences with the ESDL (and related initiatives) is to enable model development, parameterization, and evaluation. To explore this potential, we present a parameter estimation study that considers two variables only, but it helps to illustrate the approach. In fact, the approach could be extended to exploit multiple data streams in complex models. The example presented here quantifies the sensitivities of ecosystem respiration – the natural release of CO₂ by ecosystems – to fluctuations in temperature. Estimating such sensitivities is key for understanding and modelling the global climate–carbon cycle feedbacks (Kirschbaum, 1995). The following simple model (Davidson and Janssens, 2006) is widely used as a diagnostic description of this process:

\[
R_{\text{eco},i} = R_b Q_{10}^{\frac{r_i - T_{\text{ref}}}{10}},
\]

where \( R_{\text{eco},i} \) is ecosystem respiration at time point \( i \), and the parameter \( Q_{10} \) is the temperature sensitivity of this process, i.e. the factor by which \( R_{\text{eco},i} \) would change by increasing...
(or decreasing) the temperature $T_i$ by 10°C. An indication of how much respiration we would expect at some given reference temperature $T_{ref}$ is given by the pre-exponential factor $R_b$. Under this model, one can directly estimate the temperature sensitivities from some observed respiration and temperature time series. Technically, this is possible, and Eq. (16) describes a parameter estimation process as an atomic function:

$$ f_{\{\text{par}, \{\text{time} \}} : \mathcal{C}([\text{lat, long, time}]) \rightarrow \mathcal{C}([\text{lat, long, par}]) \times \mathcal{C}([\text{lat, long, time}]), \quad (16) $$

that expects a multivariate time series and returns a parameter vector. Figure 7a visualizes these estimates, which are comparable to many other examples in the literature (see, e.g. Hashimoto et al., 2015) and depict pronounced spatial gradients. High-latitude ecosystems seem to be particularly sensitive to temperature variability according to such an analysis.

However, it has been shown theoretically (Davidson and Janssens, 2006), experimentally (Sampson et al., 2007), and using model–data fusion (Migliavacca et al., 2015), that the underlying assumption of a constant base rate is not justified. The reason is that the amount of respirable carbon in the ecosystem will certainly vary with the supply, and hence phylogeny, as well as with respiration-limiting factors such as water stress (Reichstein and Beer, 2008). In other words, ignoring the seasonal time evolution of $R_b$ leads to substantially confounded parameter estimates for $Q_{10}$.

One generic solution to the problem is to exploit the variability of respiratory processes at short-term modes of variability. Specifically, one can apply a timescale-dependent parameter estimation (SCAPE; Mahecha et al., 2010b), assuming that $R_b$ varies slowly, e.g. on a seasonal and slower timescale. This approach requires some time series decomposition as described in Sect. 4.2. The SCAPE idea requires to rewrite the model, after linearization, such that it allows for a time-varying base rate:

$$ \ln R_{eco,i} = \ln R_{b,i} + \frac{T_i - T_{ref}}{10} \ln Q_{10}. \quad (17) $$

The discrete spectral decomposition into frequency bands of the log-transformed respiration allows to estimate $\ln Q_{10}$ on specific timescales that are independent of phenological state changes (for an in-depth description, see Mahecha et al., 2010b, supporting materials). Conceptually, the model estimation process now involves two steps (Eqs. 18 and 19): a spectral decomposition where we produce a data cube of higher order,

$$ f_{\{\text{par}, \{\text{time} \}} : \mathcal{C}([\text{lat, long, time}]) \rightarrow \mathcal{C}([\text{lat, long, time}, \{\text{var}, \{\text{freq} \}}]), \quad (18) $$

followed by the parameter estimation, which differs from the approach described in Eq. (16), as this approach only returns a singular parameter ($Q_{10}$), whereas $\ln R_{b,i}$ now becomes a time series:

$$ f_{\{\text{par}, \{\text{time} \}} : \mathcal{C}([\text{lat, long, time, var, freq}]) \rightarrow \mathcal{C}([\text{lat, long}]) \times \mathcal{C}([\text{lat, long, time}]). \quad (19) $$

The results of the analysis are shown in Fig. 7b, where we find generally a much more homogeneous and better constrained spatial pattern of $Q_{10}$. As suggested in the site-level analysis by Mahecha et al. (2010b) and later by others (see, e.g. Wang et al., 2018), we find a global convergence of the temperature sensitivities. We also find that, e.g. semi-arid and savanna-dominated regions clearly show lower apparent $Q_{10}$ (Fig. 7a) compared to the SCAPE approach (Fig. 7b). Discussing these patterns in detail is beyond the scope of this paper, but in general terms these findings are consistent with the expectation that in semi-arid ecosystems confounding factors act in the opposing direction (Reichstein and Beer, 2008).
4.4 Bivariate relations in vector cubes

The original idea of the data cube concept emerged from the need for working with large multivariate gridded data sets. However, the idea of data cubes can be possibly extended to other types of geographical data. One example is vector data cubes, where, e.g. polygons form an axis in their own right and each polygon points to a complex spatial shape. Consider, for instance, the need for statistical inferences on the spatial polygons often used in Intergovernmental Panel on Climate Change (IPCC) reports. One relevant question is, for example, understanding the relations of GPP and surface moisture. Figure 8 shows the bivariate histograms between both variables within a selected set of regions. This analysis clearly shows that in many regions of the world, GPP and surface moisture are strongly coupled. Examples are Cen-
tral America/Mexico (CAM), north-east Brazil (NEB), west Africa (WAF), southern Africa (SAF), east Africa (EAF), south Asia (SAS), or south Australia/New Zealand (SAU). All of these regions contain significant fractions of semi-arid climates, which can explain the constraints that water availability has on photosynthetic CO$_2$ uptake. In other regions, this relation is less obvious and often not pronounced, probably because the cases of water shortage are rare compared to the normal dynamics that might be constrained by other factors such as temperature. From a computational point of view, this example follows a very different logic, compared to the concept of applying an UDF on some of the cube axis. Rather, this example was computed using an “online” approach which sequentially updates some statistics (here the bivariate histograms) over a given class (here the IPCC regions). Such an approach allows calculations with large amounts of data and shows that the ESDL framework can also be coupled with conceptually very different analytical frameworks that might be particularly relevant when working with living data, i.e. with data streams that are constantly updated. In these cases, it is not desirable to constantly re-estimate all relevant quantities across the entire data cube.

5 Discussion

In the following, we describe the insights gained during the development of the concept and the implementation of the ESDL, addressing issues arising and critiques expressed during our community consultation processes. We also briefly discuss the ESDL in light of other developments in the field. Finally, we highlight some challenges ahead and proposed future applications.

5.1 Insights and critical perspectives

During a community consultation process across various workshops and summer schools, users expressed confusion about the equitable treatment of data cube dimensions (Sect. 2). Considering that an unordered nominal dimension of “variables” is a dimension as “time” or “latitude” seems counterintuitive at first glance. Also, concerns have been expressed about whether “time” can be treated analogously to, e.g. “latitude”. Our main argument during the development of the ESDL was that it is possible, as long as the UDFs are not applied to dimensions where they would produce non-sense results. But the practical arguments for a common interface prevail. Also, and this is key, the concept and implementation are sufficiently flexible to allow users to deploy a more classical approach to deal with such data, e.g. analysing variables separately, or writing specific UDFs that specifically require spatial or temporal dimensions. However, for research examples structured like the second use case (Sect. 4.2), the proposed approach was key, as it is allowed to efficiently navigate through the variable dimension. It is obviously irrelevant to algorithms of dimensionality reduction which dimension is compressed, and we could have equally asked the question in time domain or across a spatial dimensions, which relates to the well-known empirical orthogonal functions (EOFs) as used in climate sciences (Storch and Zwiers, 1999). In exploratory approaches of this kind, where there is no prior scientific basis for presupposing where the “information-rich zones” are in the data cube, a dimension-agnostic approach clearly pays off. We also favour this idea as it is in-line with other approaches discussed in the community. For instance, the “data cube manifesto” (Baumann, 2017) states that “data cubes shall treat all axes alike, irrespective of an axis having a spatial, temporal, or other semantics”, a principle that we have radically implemented in the ESDL.jl Julia package (Sect. 3). The flexibility we gain is that we are, in principle, prepared for comparable cases where one has to deal with, e.g. multiple model versions, model ensemble members, or model runs based on varying initial conditions.

One of the most commonly expressed practical concerns is the choice of a unique data grid. The curation of multiple data streams within such a data cube grid requires that many data have to undergo reformatting and/or remapping. Of course, this can be problematic at times, in particular when data have been produced for a given spatial or temporal resolution and cannot be remapped without violating basic assumptions. For instance, keeping mass balances, integrals of flux densities, and global moments of intensive properties as consistent as possible should always be a priority. However, for the data cube approach implemented here, we decided to accept certain simplifications. The availability of a multitude of relevant data to study Earth system dynamics is a key incentive to use the ESDL and goes far beyond many disciplinary domains. But, as we have learned in this discussion, it comes at the price of some pragmatic trade-offs. A fundamental advancement of our approach would be to natively deal with data streams from unequal grids.

The current notation of the concept has been criticized for being unsuitable for dealing with so-called vector data cubes (Pebesma and Appel, 2019). Indeed, other conceptual approaches are more suited to ours to treat such examples (see, e.g. Gebbert et al., 2019). But the research example briefly described in Sect. 4.4 and Fig. 8 does showcase such a possibility. In this case, the idea of mapping a single function across some dimensions cannot be trivially realized, but it opens novel perspectives to compute statistics based on very big data. Further research needs to be done on developing the ESDL in such directions because it would allow not only for dealing with big data issues but also to update statistics without having to recompute data processed in earlier steps. This can solve the challenges of dealing with “living data”.

One of the main concerns expressed by users, in particular by 30 young researchers who participated in the project during an early adopter phase, is the demand for the latest data in the ESDL. This is why the concept presented here and its implementation should be further developed into a persis-
tent infrastructure. Such a step is challenging and there is a trade-off to be made between wishing to include latest data streams (ideally even in near-real time) and constantly expanding the access API and portfolio of example workflows. The ESDL thus depends on the enduring enthusiasm of the user community and funding agencies to support the idea in this respect and grow steadily into new domains, help us add data streams, and actively co-develop the approach.

5.2 Relation to other initiatives and platforms

Over the past few years, several initiatives, platforms, and software solutions (Lu et al., 2018; Sudmanns et al., 2019) have emerged based on similar considerations as those motivating the Earth System Data Lab. Some of these platforms and software solutions are explicitly constructed around the idea of data cubes (e.g. Baumann et al., 2016; Lewis et al., 2017; Appel and Pebesma, 2019). Nevertheless, the concept of “data cube” is still not fully consolidated in the Earth system science. It was only in 2019 that the Open Geospatial Consortium (OGC) opened a public discussion towards establishing standards for data cubes.

Among the other existing initiatives, the Climate Data Store (CDS) of the Copernicus Climate Change Service (https://cds.climate.copernicus.eu/, last access: 21 February 2020) is conceptually probably the closest one to the ESDL. The CDS was primarily designed as key infrastructure to analyse climate reanalysis data and related variables. These data often require to be analysed at very high temporal resolutions (e.g. using hourly time steps). The CDS offers a similar Python interface to analyse these data. Likewise, the Google Earth Engine (GEE; https://earthengine.google.com, last access: 21 February 2020; Gorelick et al., 2017) is probably the most widely known platform for implementing global-scale analytics. GEE offers access to a wide range of satellite data archives and increasingly also to climate data in their native resolutions. One strength of GEE is the massive computing power offered to the scientist, such that some use cases nicely showcased the power of the infrastructure. The user has a wide range of predefined operators available that can be used and coupled to build workflows that are particularly suitable for time series. Another recent development in the field is the Open Data Cube (ODC; https://www.opendatacube.org/, last access: 21 February 2020; formerly Australian Data Cube; Lewis et al., 2017). This project was initially designed to offer access to the well-processed remote sensing data over Australia with an emphasis on the Landsat archive. In the past years, the ODC technology was used to implement regional data cubes for Colombia (CDCol; Ariza-Porras et al., 2017; Bravo et al., 2017), Switzerland (SDC; http://www.swissdatacube.org/, last access: 21 February 2020; Giuliani et al., 2017), and Armenia (Asmryan et al., 2019), among many other countries. The aim of the open-access ODC is also to effectively enable access to time series data from high-resolution data archives, targeting mainly changes in land surface properties. The ESDL has developed into a conceptually different direction than most of the other initiatives that make it unique.

First, we note that most of the data cube initiatives were motivated by the need to access and/or analyse big, e.g. very-high-resolution, data (Lewis et al., 2017; Nativi et al., 2017; Giuliani et al., 2019). Initially, this problem was not in the focus of the ESDL, which rather aimed at downstream data products. Our data cube approach primarily intends to support the joint exploitation of multiple data streams efficiently. This multivariate focus is rarely found as a key design element in the other approaches.

Second, most initiatives intend to preserve the resolutions of the underlying data. The ESDL, instead, is built around singular data cubes that then include variables as an additional dimension. The inevitable trade-off, as discussed above, is the need for a data curation and remapping process prior to the analyses.

Third, there is a wide consensus that data cube technologies need to enable the application of UDFs. However, at this stage, this aspect often appears not to be a priority of other data cube initiatives and, consequently, users are restricted in their analysis by the available tools. In this context, we see the strength of the ESDL, as it allows for the development of complex workflows and adding arbitrary functionalities efficiently. This is actually one reason why we decided to implement the ESDL in the quite young language of scientific computing Julia (side by side with the more commonly used Python tools).

Taken together, the ESDL has probably conceptually developed (and implemented) the most radical cubing principle following a strict dimension agnostic approach. We envisage that the ESDL front end could be coupled to a data cube technology as proposed by any of the other initiatives to combine its analytic strength with the efficiencies achieved by others in dealing with high-resolution data streams.

5.3 Priorities for future developments

During the development of the ESDL, we identified several methodological challenges on the one hand and, on the other, application domains that could be addressed. With regard to potentially relevant methodological paths, we can only briefly mention, with no claim to completeness, some of the most ardently and widely discussed topics:

- **Machine learning.** Data-driven approaches have always been part of the DNA of Earth system sciences (see classical textbooks, e.g. Storch and Zwiers, 1999) and classically complement process-driven modelling efforts (Luo et al., 2012). However, with the rise of modern machine learning, new perspectives have emerged (Mjolness and DeCoste, 2001; Hsieh, 2009). Depending on the purpose, we find purely exploratory analysis based on, e.g. nonlinear dimensionality reduction (Mahecha
Spatial interactions. For interpreting the interactions and mechanisms of the land and ocean, or land and atmosphere that involve lateral transport, the ESDL would require more developments. Statistical approaches like spatial network analyses (e.g., Donges et al., 2009; Boers et al., 2019) or process-oriented ideas like explicit moisture transport (e.g., Wang-Erlandsson et al., 2018) would be very valuable to be explored but would require a substantial rethinking of the actual implementation in order to achieve high performance.

Model evaluation and benchmarking. Our third use case (Sect. 4.3) illustrates the suitability of the ESDL for parameter estimation and model evaluation purposes. Today, typical model evaluation frameworks in the Earth system sciences prepare predefined benchmark metrics on some reference data sets (Luo et al., 2012). Prominent examples are the benchmarking tools awaiting the sixth phase of the Coupled Model Intercomparison Project (CMIP6) model suites (Eyring et al., 2019). However, these model evaluation frameworks typically do not give the user the full flexibility to apply some user-defined metrics to the model ensemble under scrutiny. We believe that mapping UDFs on such big Earth system model output could greatly benefit the development of novel evaluation metrics in the near future. Building data cubes from multi-model ensembles would be straightforward, as different models or ensembles would simply lead to one additional dimension in our setup. In fact, the ESDL approach is perfectly suited to handle, e.g., the output of the actual CMIP data, as we have already exemplified. Of course, any other model ensembles can be treated analogously.

In terms of application domains, we see high potential in the following areas:

Human–environment interactions. Addressing the complexities of human–environment interactions (Schimel et al., 2015) is a particular challenge. Making the ESDL fit for this purpose would require integrating a variety of (at least) spatially explicit population estimates (Doxsey-Whitfield et al., 2015) and socioeconomic data (Smits and Permanyer, 2019). The latter represent a fundamentally novel development that has great potential for understanding, e.g., dynamics of disaster impacts (Guha-Sapir and Checchi, 2018), among other issues. In fact, this integration is a grand challenge ahead (Mahecha et al., 2019) but not out of reach for the ESDL.

Biodiversity research. Another question of high societal relevance is to understand how patterns of biodiversity affect ecosystem functioning (Emmott Duffy et al., 2017; García-Palacios et al., 2018). In light of a global decline in species richness (see latest global reports; https://www.ipbes.net/, last access: 21 February 2020), this question is of uttermost importance. The ESDL is only partly fit for this purpose, as it would require the ingestion of a wide range of essential biodiversity variables (Pereira et al., 2013; Skidmore et al., 2015), beyond the ones we have already available. But still, the ESDL is conceptually prepared to deal with these challenges (compare, e.g., the demands described in Hardisty et al., 2019) and would be particularly suitable for relating biodiversity patterns to the so-called ecosystem function properties (Reichstein et al., 2014; Musavi et al., 2015). In fact, in the regional application of the ESDL, we have focused on Colombia and its wider region to explore linkages of this kind relying on remote-sensing-derived variables that are relevant for this context.

Oceanic sciences. Extending the ESDL for ocean data is desired and conceptually possible. Surface parameters, e.g., phytoplankton phenology derived from remote sensing (Racault et al., 2012), can be treated analogously to terrestrial surface parameters. Other dynamics, e.g., the analysis and exploration of ocean–land coupling mechanisms, ocean–atmosphere interactions, and land–atmosphere interactions triggered by ocean circulation dynamics, could in principle be facilitated via the ESDL but require either vertical or lateral dynamics.

Solid Earth. The step towards global, fully data informed model data is also made in geophysics. For
instance, recently Afonso et al. (2019) used an inversion approach to develop a 3-D model that fully describes multiple parameters in the Earth interior, including, e.g. crustal and lithospheric thickness, average crustal density, and a depth-dependent density of the lithospheric mantle, among other variables. They proposed a tool allowing for inspecting the data interactively at a spatial resolution of $2^\circ \times 2^\circ$ grid at different depths. Clearly, in this case, other dimensions are relevant, but the principle remains the same and, in fact, can be treated in a very similar manner. Future model–data assimilation approaches of this kind could be performed in the context of the ESDL, as well as the aforementioned machine learning for the solid Earth (Bergen et al., 2019).

In summary, we have demonstrated that the ESDL is a flexible and generic framework that can allow various different communities to explore and analyse large amounts of gridded data efficiently. Thinking about the potential paths ahead, the ESDL could become a valuable tool in various fields of Earth system sciences, biodiversity research, computer sciences, and other branches of science. The widespread social and political uptake of the concept of planetary boundaries (Rockström et al., 2009; Steffen et al., 2015) underlines the global demand for better quantified process understanding of environmental risks and resource bottlenecks based on empirical evidence. Along these lines, the ESDL concept could be used to address some of the most pressing global challenges. For example, it could become an interface for direct interaction with ECVs, global climate projections, and EBVs. Such an interactive interface would allow a much broader community to better understand the data underlying the global assessment reports of the IPCC (IPCC, 2014) and Intergovernmental Science-Policy Platform on Biodiversity and Ecosystem Services (IPBES) (Díaz et al., 2019). If coupled to some visual interfaces, the ESDL could also be used by a broader community, enhancing education, communication, and decision-making process, contributing to knowledge democratization about a deeper understanding of the complex and dynamic interactions in the Earth system.

6 Conclusions

Exploiting the synergistic potential of multiple data streams in the Earth sciences beyond disciplinary boundaries requires a common framework to treat multiple data dimensions, such as spatial, temporal, variable, frequency, and other grids alike. This idea leads to a data cube concept that opens novel avenues to efficiently deal with data in the Earth system sciences. In this paper, we have formalized the concept of data cubes and described a way to operate on them. The outlined dimension-agnostic approach is implemented in the Earth System Data Lab, which enables users applying a wide range of functions to all thinkable combinations of dimension. We believe that this idea can dramatically reduce the barrier to exploit Earth system data and serves multiple research purposes. The ESDL complements a range of emerging initiatives that differ in architectures and specific purposes. However, the ESDL is probably the most radical data cubing approach, offering novel opportunities for cross-community data-intensive exploration of contemporary global environmental changes. Future developments in related branches of science and latest methodological developments need to be considered and addressed soon. At its actual state of implementation, the ESDL can already contribute to the deeper understanding and more effective implementation of policy-relevant concepts such as the planetary boundaries, essential variables in different subsystems of the Earth, and global assessment reports. We see a particularly high future potential for data cube concepts as presented for, firstly, interpreting large-scale model ensembles, and secondly, analysing new multispectral satellite remote sensing data with their constantly increasing spatial, temporal, and spectral resolutions.
Appendix A: Data streams in the Earth System Data Lab

In the following, we give an overview of the actually available variables in the Earth System Data Lab. The list is constantly being updated.
Table A1. Data streams in the current implementation of the ESDL.

<table>
<thead>
<tr>
<th>Domain</th>
<th>Variable</th>
<th>Short</th>
<th>Coverage</th>
<th>Description</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Atmosphere</td>
<td>2 m temperature</td>
<td>$T_{2m}$</td>
<td>2001–2011</td>
<td>The 2 m air temperature data ($[T_{2m}] = \text{K}$) are part of the ERA-Interim reanalysis product and therefore produced by data assimilation techniques in combination with a forecast model. The original spatial sampling (T255 spectral resolution) approximates to 80 km and the original temporal sampling is 6 h for analyses and 3 h for forecasts.</td>
<td>Dee et al. (2011)</td>
</tr>
<tr>
<td>Atmosphere</td>
<td>Aerosol optical thickness at 550 nm</td>
<td>AOD$_{550}$</td>
<td>2002–2012</td>
<td>The ESA CCI aerosol optical thickness (depth) data sets were created by using algorithms which were developed in the ESA aerosol_cci project. The data used here were created from Advanced Along-Track Scanning Radiometer (AATSR) measurements (ENVISAT mission) using the algorithm and represent total column AOD at the specified wavelength. Horizontal resolution of the daily data is $1^\circ \times 1^\circ$ on a global grid.</td>
<td>Holzer-Popp et al. (2013)</td>
</tr>
<tr>
<td>Atmosphere</td>
<td>Aerosol optical thickness at 555 nm</td>
<td>AOD$_{555}$</td>
<td>2002–2012</td>
<td>The ESA CCI aerosol optical thickness (depth) data sets were created by using algorithms which were developed in the ESA aerosol_CCI project. The data used here were created from AATSR measurements (ENVISAT mission) using the... algorithm and represent total column AOD at the specified wavelength. Horizontal resolution of the daily data is $1^\circ \times 1^\circ$ on a global grid.</td>
<td>Holzer-Popp et al. (2013)</td>
</tr>
<tr>
<td>Atmosphere</td>
<td>Aerosol optical thickness at 659 nm</td>
<td>AOD$_{659}$</td>
<td>2002–2012</td>
<td>The ESA CCI aerosol optical thickness (depth) data sets were created by using algorithms which were developed in the ESA aerosol_cci project. The data used here were created from AATSR measurements (ENVISAT mission) using the... algorithm and represent total column AOD at the specified wavelength. Horizontal resolution of the daily data is $1^\circ \times 1^\circ$ on a global grid.</td>
<td>Holzer-Popp et al. (2013)</td>
</tr>
<tr>
<td>Atmosphere</td>
<td>Aerosol optical thickness at 865 nm</td>
<td>AOD$_{865}$</td>
<td>2002–2012</td>
<td>The ESA CCI aerosol optical thickness (depth) data sets were created by using algorithms which were developed in the ESA aerosol_cci project. The data used here were created from AATSR measurements (ENVISAT mission) using the... algorithm and represent total column AOD at the specified wavelength. Horizontal resolution of the daily data is $1^\circ \times 1^\circ$ on a global grid.</td>
<td>Holzer-Popp et al. (2013)</td>
</tr>
<tr>
<td>Atmosphere</td>
<td>Aerosol optical thickness at 1610 nm</td>
<td>AOD$_{1610}$</td>
<td>2002–2012</td>
<td>The ESA CCI aerosol optical thickness (depth) data sets were created by using algorithms which were developed in the ESA aerosol_cci project. The data used here were created from AATSR measurements (ENVISAT mission) using the... algorithm and represent total column AOD at the specified wavelength. Horizontal resolution of the daily data is $1^\circ \times 1^\circ$ on a global grid.</td>
<td>Holzer-Popp et al. (2013)</td>
</tr>
</tbody>
</table>
Table A1. Continued.

<table>
<thead>
<tr>
<th>Domain</th>
<th>Variable</th>
<th>Short</th>
<th>Coverage</th>
<th>Description</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Biosphere</td>
<td>Gross primary productivity</td>
<td>GPP</td>
<td>2001–2012</td>
<td>By training an ensemble of machine learning algorithms with eddy covariance data from FLUXNET and satellite observations in a cross-validation approach, regressions from these observations to different kinds of carbon and energy fluxes were established and used to generate data sets with a spatial resolution of 5 arcmin and a temporal resolution of 8 d. The GPP resembles the total carbon release of the ecosystem through respiration and is expressed in units of gC m$^{-2}$ d$^{-1}$.</td>
<td>Tramontana et al. (2016)</td>
</tr>
<tr>
<td>Biosphere</td>
<td>Net ecosystem exchange</td>
<td>NEE</td>
<td>2001–2012</td>
<td>By training an ensemble of machine learning algorithms with eddy covariance data from FLUXNET and satellite observations in a cross-validation approach, regressions from these observations to different kinds of carbon and energy fluxes were established and used to generate data sets with a spatial resolution of 5 arcmin and a temporal resolution of 8 d. The NEE resembles the net carbon exchange between the ecosystem and the atmosphere and is expressed in units of gC m$^{-2}$ d$^{-1}$.</td>
<td>Tramontana et al. (2016)</td>
</tr>
<tr>
<td>Land</td>
<td>Latent energy</td>
<td>$LE$</td>
<td>2001–2012</td>
<td>By training an ensemble of machine learning algorithms with eddy covariance data from FLUXNET and satellite observations in a cross-validation approach, regressions from these observations to different kinds of carbon and energy fluxes were established and used to generate data sets with a spatial resolution of 5 arcmin and a temporal resolution of 8 d. The $LE$ resembles the latent heat flux from the surface and is expressed in units W m$^{-2}$.</td>
<td>Tramontana et al. (2016)</td>
</tr>
<tr>
<td>Land</td>
<td>Sensible heat</td>
<td>$H$</td>
<td>2001–2012</td>
<td>By training an ensemble of machine learning algorithms with eddy covariance data from FLUXNET and satellite observations in a cross-validation approach, regressions from these observations to different kinds of carbon and energy fluxes were established and used to generate data sets with a spatial resolution of 5 arcmin and a temporal resolution of 8 d. The $H$ resembles the sensible heat flux from the surface and is expressed in units W m$^{-2}$.</td>
<td>Tramontana et al. (2016)</td>
</tr>
<tr>
<td>Land</td>
<td>Monthly burnt area</td>
<td>Burnt area</td>
<td>1995–2014</td>
<td>This data set was taken from the fourth generation of the Global Fire Emissions Database (GFED4). It was created as a combination of data from infrared sensor satellite observations and resembles the estimated monthly burnt area in hectares. The spatial resolution of this data set is 0.25°. Small fires were exempt in the production of the data.</td>
<td>Giglio et al. (2013)</td>
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<tr>
<td>Domain</td>
<td>Variable</td>
<td>Short</td>
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<tr>
<td>Land</td>
<td>Carbon dioxide emissions due to natural fires expressed as carbon flux</td>
<td>Emission</td>
<td>2001–2010</td>
<td>This data set was taken from the fourth generation of the Global Fire Emissions Database (GFED4). It was created by applying a model based on the Carnegie–Ames–Stanford approach (CASA) to the burnt area estimates and has the same temporal (monthly) and spatial (0.25°) resolution as the monthly burnt area data set and expresses the carbon dioxide emissions of natural fires as a carbon flux (gC m⁻² d⁻¹). Small fires were included in this approach.</td>
<td>Giglio et al. (2013), van der Werf et al. (2017)</td>
</tr>
<tr>
<td>Land</td>
<td>Evaporation</td>
<td>E</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semiempirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>Evaporative stress factor</td>
<td>S</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semiempirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
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<tr>
<td>Land</td>
<td>Potential evaporation</td>
<td>E_p</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semiempirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>Interception loss</td>
<td>E_i</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semiempirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
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<td>Variable</td>
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<tr>
<td>Land</td>
<td>Root-zone soil moisture</td>
<td>SM&lt;sub&gt;root&lt;/sub&gt;</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semi-empirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>Surface soil moisture</td>
<td>SM&lt;sub&gt;surf&lt;/sub&gt;</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semi-empirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>Bare soil evaporation</td>
<td>E&lt;sub&gt;b&lt;/sub&gt;</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semi-empirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>Snow sublimation</td>
<td>E&lt;sub&gt;s&lt;/sub&gt;</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellites, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semi-empirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>Transpiration</td>
<td>E&lt;sub&gt;t&lt;/sub&gt;</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellite sensors, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semi-empirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
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<td>Domain</td>
<td>Variable</td>
<td>Short</td>
<td>Coverage</td>
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<tr>
<td>Land</td>
<td>Open-water evaporation</td>
<td>$E_w$</td>
<td>2001–2011</td>
<td>The GLEAM data sets are created by using a set of algorithms, input forcing data sets from reanalyses, optical and microwave satellite sensors, and other merged sources. The model itself consists of four modules: potential evaporation (Priestley–Taylor equation), interception (Gash analytical model), soil (multilayer soil model plus data assimilation), and stress (semi-empirical). The data are sampled on a graticule of 0.25° and have a daily temporal coverage.</td>
<td>Martens et al. (2017), Miralles et al. (2011)</td>
</tr>
<tr>
<td>Land</td>
<td>White-sky albedo for visible wavelengths</td>
<td>BHR_VIS</td>
<td>1998–2012</td>
<td>White-sky albedo, also known as bihemispherical reflectance (only diffuse illumination), estimated from satellite radiometer data. The spatial resolution of this product is 1 km with a temporal sampling of 8 d.</td>
<td>Lewis et al. (2012)</td>
</tr>
<tr>
<td>Land</td>
<td>Black-sky albedo for visible wavelengths</td>
<td>DHR_VIS</td>
<td>1998–2012</td>
<td>Black-sky albedo, also known as directional–hemispherical reflectance (only direct illumination), estimated from satellite radiometer data. The spatial resolution of this product is 1 km with a temporal sampling of 8 d.</td>
<td>Lewis et al. (2012)</td>
</tr>
<tr>
<td>Water</td>
<td>Fractional snow cover</td>
<td>MFSC</td>
<td>2003–2013</td>
<td>Global fractional snow cover product using mainly satellite infrared radiometer data (ATSR-2, AATSR). Glaciers, continental ice shields, and snow on ice are exempt from the data. Values stand for the percentage of the area of a grid cell covered by snow integrated over time (daily, weekly, or monthly). The spatial resolution is 1 km.</td>
<td>Luojus et al. (2010), Metsämäki et al. (2015)</td>
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<tr>
<td>Water</td>
<td>Snow water equivalent</td>
<td>SWE</td>
<td>1980–2012</td>
<td>Snow water equivalent product covering the Northern Hemisphere (35–85° N), created by using microwave sensor data (SMMR, SSM/I, SSMIS). Glaciers, continental ice shields, and mountainous regions are exempt from the data. Values stand for the water equivalent of snow per grid cell in millimetres aggregated over time (daily, weekly, or monthly). The weekly data are produced by giving every day the mean value of a sliding window (−6 d). The monthly data are given as the weekly mean and maximum per calendar month. The spatial resolution is approximately 25 km.</td>
<td>Luojus et al. (2010)</td>
</tr>
<tr>
<td>Land</td>
<td>Land surface temperature</td>
<td>LST</td>
<td>2002–2011</td>
<td>The GlobTemperature Land Surface Temperature product used here is a product of a satellite infrared radiometer (AATSR). It has global coverage with a spatial sampling of 0.05° and consists of two measurement averages (day and night). The values are an approximation of the average land surface temperature per grid cell in K. It is an improved version of the ESA AATSR data set (UOL_LST_3P, v2.1).</td>
<td>Ghent (2012)</td>
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<tr>
<td>Domain</td>
<td>Variable</td>
<td>Short</td>
<td>Coverage</td>
<td>Description</td>
<td>References</td>
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<tr>
<td>Atmosphere</td>
<td>Total column water vapour</td>
<td>TCWV</td>
<td>1996–2008</td>
<td>The TCWV product was derived through combination of various satellite spectrometer and microwave sensor data sets. It resembles the total mass of water contained in a column of air from the surface to 200 hPa. The unit is kg m(^{-2}), the spatial sampling is 0.5(^{\circ}), and the data are provided as daily composites. From 1996 to 2002 (inclusive), the data consist of weekly/monthly means.</td>
<td>Schröder et al. (2012), Schneider et al. (2013)</td>
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<tr>
<td>Atmosphere</td>
<td>Mean total ozone column</td>
<td>Ozone</td>
<td>1996–2011</td>
<td>The total ozone column data from the Ozone CCI project is derived from the Global Ozone Monitoring Experiment (GOME) spectrometer acquisitions. For the ESDL, level 2 data have been used. They are given in Dobson units (DU) and have a spatial resolution of 320 km x 40 km. The temporal resolution depends on the latitude, with the longest revisit time being 3 d at the Equator.</td>
<td>Van Roozendael et al. (2012), Lerot et al. (2014)</td>
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<tr>
<td>Land</td>
<td>Fraction of absorbed photosynthetically active radiation</td>
<td>fAPAR_TIP</td>
<td>1982–2016</td>
<td>The fAPAR, describing the amount and productivity of vegetation, was derived by using a two-stream inversion package (TIP) method based on the two-stream model developed by Pinty et al. (2006). The product is delivered in two spatial resolutions (0.05 and 0.5(^{\circ})) and with a daily temporal coverage.</td>
<td>Disney et al. (2016), Blessing and Löw (2017)</td>
</tr>
<tr>
<td>Land</td>
<td>Leaf area index</td>
<td>LAI</td>
<td>1982–2016</td>
<td>The LAI, defined as half the total canopy area per unit ground area (m(^2) m(^{-2})), was derived by using a TIP method based on the two-stream model developed by Pinty et al. (2006). The product is delivered in two spatial resolutions (0.05 and 0.5(^{\circ})) and with a daily temporal coverage.</td>
<td>Disney et al. (2016), Blessing and Löw (2017)</td>
</tr>
<tr>
<td>Land</td>
<td>White-sky albedo for visible wavelengths from AVHRR</td>
<td>BHR_VIS</td>
<td>1982–2016</td>
<td>White-sky albedo, also known as bihemispherical reflectance (only diffuse illumination), estimated from satellite radiometer data. This data set extends the GlobAlbedo data by using additional input data sources (Advanced Very High Resolution Radiometer (AVHRR), geostationary satellites). The product is delivered in two spatial resolutions (0.05 and 0.5(^{\circ})) and with a daily temporal coverage.</td>
<td>Lewis et al. (2012), Danne et al. (2017)</td>
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<tr>
<td>Land</td>
<td>Black-sky albedo for visible wavelengths from AVHRR</td>
<td>DHR_VIS</td>
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<td>Black-sky albedo, also known as directional–hemispherical reflectance (only direct illumination), estimated from satellite radiometer data. This data set extends the GlobAlbedo data by using additional input data sources (AVHRR, geostationary satellites). The product is delivered in two spatial resolutions (0.05 and 0.5(^{\circ})) and with a daily temporal coverage.</td>
<td>Lewis et al. (2012), Danne et al. (2017)</td>
</tr>
<tr>
<td>Land</td>
<td>Fraction of absorbed photosynthetically active radiation from AVHRR</td>
<td>fAPAR_AVHRR</td>
<td>1982–2006</td>
<td>The AVHRR-derived fAPAR, describing the amount and productivity of vegetation, was derived from AVHRR black-sky albedo data. The product is delivered in two spatial resolutions (0.05 and 0.5(^{\circ})) and with a daily temporal coverage.</td>
<td>Gobron et al. (2017)</td>
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## Table A1. Continued.

<table>
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<tr>
<th>Domain</th>
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<th>Short</th>
<th>Coverage</th>
<th>Description</th>
<th>References</th>
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<tbody>
<tr>
<td>Land</td>
<td>Soil moisture</td>
<td>SM</td>
<td>1978–2017</td>
<td>The ESA CCI soil moisture data combine various active and passive microwave sensors into a homogenized product. It represents the soil water content in the upper 5 cm of the soil, produced at a spatial sampling of 0.25° and a temporal sampling of 1 d. Gaps exist in periods of snow cover or frozen conditions and in areas with very dense vegetation.</td>
<td>Liu et al. (2012), Dorigo et al. (2017), Gruber et al. (2017)</td>
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</table>
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**Data availability.** All data are available via https://www.earthsystemdatalab.net/ (last access: 21 February 2020) or from the original data providers as indicated in Table A1 in the paper.

**Author contributions.** MDM, FG, and MR developed the concept; FG implemented the ESDL.jl package in the Julia language; NF, FG, MDM, and GB implemented the overall project; MCL and LES implemented the Colombian cube. RC, JP, MDM, FG, GK, PB, and PP developed the notation; DGM and MJ contributed particular data. MDM wrote the manuscript with substantial input from FG, SC, GCV, RC, JP, and GK, and detailed comments from all co-authors.

**Competing interests.** The authors declare that they have no conflict of interest.

**References**


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