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1 **Modelling water balance and nitrate leaching in temperate Norway**
2 **spruce and beech forests located on the same soil type with the**
3 **CoupModel**

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26 **Abstract**

27 Two contrasting forest ecosystems located in close proximity to each other were selected for evaluating the
28 importance of tree species and afforestation in relation to the water balance and the quality of the water
29 leaving the forest root zone. Measurements included soil water content and the collection of precipitation,
30 canopy throughfall, stem flow and soil solution on a weekly basis during 15 months (1999–2000). Soil
31 solutions were extracted using suction probes installed at all major horizons within the upper 120 cm of a
32 Norway spruce (N. spruce) stand (Picea Abies [L.] Karst.) and a European beech stand (Fagus Sylvatica L.)
33 located on the same soil type. Soil solutions were analyzed for the content of all major ions, including nitrate.
34 A water balance model (CoupModel) was used to estimate percolation rates beneath the root zone.
35 Percolation at the beech stand was 292 mm and only 41 mm at the N. spruce stand mainly due to differences
36 in the interception loss. The highest annual leaching of Mg, K, Na, Al, Cl, SO₄-S was noted in the N. spruce
37 stand while leaching of NO₃-N was highest in the beech stand, corresponding to 39 kg ha⁻¹ y⁻¹. By contrast,
38 the annual leaching of NO₃-N in the N. spruce stand was only 0.5 kg ha⁻¹ y⁻¹. The larger amount of NO₃-N
39 was leaving the beech forest soil despite the fact that the N. spruce stand had the highest atmospheric N-
40 deposition. Thus, differences in NO₃-N leaching between the stands must be related to differences in uptake
41 and accumulation of N in the vegetation and within the upper 120 cm of the soil. Differences in the water
42 balance and NO₃-N leaching between beech and N. spruce stands call for further attention to the selection of
43 tree-species on a soil type basis when planning future afforestation projects, particularly when such projects
44 aim to improve the quality of water infiltrating to the groundwater zone.

45

46 Keywords: CoupModel, European beech, forest, nitrate, Norway spruce, water balance

47 **1. Introduction**

48 The number of afforestation projects is increasing due to the fact that forests can accumulate and store
49 atmospheric carbon in biomass and due to reduced leaching of nitrate as forest ecosystems are often less
50 fertilized as compared to conventional farming. In Denmark, the aim of afforestation of former agricultural
51 land is mainly to protect groundwater resources, create recreational areas and establish green corridors for
52 wildlife (Skov og Naturstyrelsen, 1999).

53 Input and turnover of nutrients in forests control the quality of percolating water. Nitrogen (N) has been
54 intensely studied due to its dual role as a vital nutrient for vegetation and as a contaminant in groundwater. N
55 circulation in particular is closely related to the cycling of carbon, due to the fact that almost all N in the
56 forest ecosystem is stored in organic form and in the same pools as carbon. Anthropogenic inputs of
57 atmospheric N over the last 30-40 years in European temperate forests may have led to a decline in forest
58 growth and elevated levels of NO₃-N leaching from forests caused by N-saturation of the ecosystem (Aber,
59 1992; Aber *et al.*, 1998; Callesen *et al.*, 1999; Dise & Wright, 1995; Gundersen, 1991; Nihlgård, 1985).

60 Groundwater recharge is controlled by the combination of atmospheric, soil and plant related processes.
61 Deciduous forest ecosystems generally yield more water and of better quality as groundwater than coniferous
62 forest ecosystems due to the smaller atmospheric deposition in the canopy (Hansen, 2003). But factors such
63 as stand age and plot can have an adverse effect on the composition of soil water (Callesen *et al.*, 1999).
64 Despite the fact that afforestation of former agricultural land will alter the hydrological cycle and water
65 balance, knowledge about changes of the amount of percolation and the quality of the soil water in forests
66 due to afforestation is scarce. This is partly due to the fact that quantitative estimates of evapotranspiration is
67 technically complicated and associated with uncertainty in measurement procedure (Wilson *et al.*, 2001).

68 One way to quantify the constituents of the water balance in forest ecosystems is to use water balance
69 models based on soil, vegetation and atmosphere characteristics (SVAT-models). These models consider the
70 interaction between meteorology, vegetation and the soil and may after acceptable calibration and validation
71 produce outputs regarding evapotranspiration, percolation and other variables difficult to measure in the
72 field. Another advantage of SVATs is that different types of vegetation can be compared under the same

73 climatic conditions. The CoupModel is a well-established SVAT-model (Jansson & Halldin, 1979), which
74 has been revised several times since (Jansson & Karlberg, 2004). The CoupModel was originally developed
75 for Swedish forest ecosystems, and has been developed to encompass most types of ecosystems. For a
76 comprehensive list of works including usage or descriptions of the CoupModel, see Jansson Karlberg (2004).
77 Recently, Ladekarl et al. (2005) used the CoupModel to calculate the water balance in oak, heath and
78 agricultural ecosystems. Measurements of the water content of the soil provide a basis for evaluating the
79 performance of the model and estimating the water balance of different ecosystems that have well defined
80 boundary conditions (Alvenäs & Jansson, 1997; Bouten & Jansson, 1995; Eckersten et al.; 1995, Jansson et
81 al.; 1999a Ladekarl et al.; 2005).

82 The main aims of this paper are firstly to quantify the water balance, using the CoupModel, in two
83 contrasting forest ecosystems, N. spruce (Picea Abies [L.] Karst.) and European beech (Fagus Sylvatica L.)
84 located on the same soil type, and secondly, to quantify and discuss the total and seasonal trends of the
85 leaching of cations and anions with a focus on NO₃-N from the two forest ecosystems.

86

87 **2. Materials & Methods**

88 2.1. Site description

89 The study site is located near Nødebo in the northern part of Zealand (55°N, 12°E), Denmark, and is
90 described in detail by Elberling & Ladegaard-Pedersen (2005). Two contrasting forest stands located within
91 2 ha of each other were selected: deciduous beech (Fagus Sylvatica L.) and coniferous common or N. spruce
92 (Picea Abies [L.] Karst.) (Fig. 1). The beech forest stand was planted in 1977 and has been growing at a rate
93 of 8 – 9 m³ ha⁻¹ yr⁻¹ to an average height of 9 m (in 2003). The N. spruce stand was planted in 1959 and has
94 been growing at a rate of 16 – 17 m³ ha⁻¹ yr⁻¹ to an average height of 23 m (in 2003). The number of trees and
95 current live aboveground volume of beech and N. spruce wood has been estimated to be approx. 600 trees ha⁻¹
96 and 145 m³ ha⁻¹ and approx. 300 trees ha⁻¹ and 387 m³ ha⁻¹, respectively (personal comm. with forest ranger
97 S. Løw, 2003). The forest is a production forest where 10 – 20% of the above ground biomass is cut every 4
98 – 5 years. Under-storey vegetation is scarce in both forest stands. No fertilizer has been used.

99 The forest stands are situated on ice marginal hills and sandy ground till from the Weichsel ice age with
100 very low or no slope in the study area. Soils have been classified as Typic Udorthents according to Soil
101 Taxonomy (Soil Survey Staff, 1997), the texture is predominantly loamy sand, and stones of varying sizes
102 are present in the soil. The main soil characteristics are presented in table 1.

103 The climate is temperate humid with a mean annual temperature of 8.2 °C (1961-1990). The mean annual
104 precipitation was approx. 657 mm with an annual potential evapotranspiration of approx. 571 mm (DMI,
105 2000). Values were calculated by DMI using interpolation algorithms established for 10 km x 10 km grids
106 covering the whole country. Algorithms were based on measurements (1961 – 1990) for meteorological
107 stations distributed evenly in Denmark. Meteorological measurements at the study site were conducted from
108 the beginning of November 1999 until the beginning of February 2001.

109

110 2.2. Measurements of precipitation, throughfall and stemflow

111 Precipitation was collected in a plastic funnel with an area of 213 cm² placed 35 cm above the ground in a
112 clearing. The sampling containers were placed in pits to limit biological activity. Throughfall (TF) was
113 collected in two plastic funnels in each stand with an area of 213 cm² and connected to an open plastic
114 container by a tube. The TF containers were also placed in a pit. Filters were put over the top of the
115 containers to avoid leaves and animals contaminating the sample. The stemflow (SF) was collected through a
116 1 m long PVC tube wrapped 1.5 times around the trunks of two separate trees in each stand and sealed with
117 silicone along the trunk to avoid water running beneath the tube. Holes of 5 mm were drilled at 8 cm
118 intervals in each tube and equipped with water filters at the beginning. The tubes were connected to a closed
119 bucket on the ground. Water sampling occurred weekly, but every second week in January and February
120 2000. The sampling frequency was changed to every third week from June 2000.

121

122 2.3. Measurements of soil water content

123 Soil water content was measured using TDR (Time Domain Reflectometry) using the approach suggested
124 by Topp et al. (1980). In November 1999 three sets of TDR-probes, 20, 40, 60, 80 and 120 cm in length and

125 5 mm in diameter, were installed in each tree stand vertically in the mineral soil after removal of the O-
126 horizon. After installation of the probes, the O-horizon material was carefully put back into place. Thus, soil
127 water content measurements represented integrated measurements in the intervals 0-20, 0-40, 0-60, 0-80 and
128 0-120 cm. Measurements were made with a Tektronix 1502C cable tester (Tektronix, 1999). Subsequently,
129 data was processed with AUTOTDR (Prenart Equipment Aps, Denmark) to estimate water content in volume
130 percentage. Water content was measured at least weekly in both stands from November 6, 1999 to February
131 16, 2001. Afterwards, the water content measurement was converted to depth-specific water storage in mm.

132

133 2.4. Collection of soil samples and soil solution

134 Intact depth-specific and volume-specific (100 cm^3) soil samples (3 replicates) were collected in October
135 2000 at 5 cm depth intervals within each horizon to a depth of 1 m. Soil samples were kept cold and dark
136 until analyzed.

137 Soil solution was extracted using teflon lysimeters of the type PRENART SUPER QUARTS (Prenart
138 Equipment Aps, Denmark) with a pore size of $2 \mu\text{m}$. One lysimeter in each horizon was installed using slurry
139 of the horizon specific soil and double ion-exchanged water in November 1999. In the beech stand the
140 lysimeter was installed at the following depths: 8, 17, 30, 48 and 70 cm. In the N. spruce stand the lysimeter
141 was installed at depths of 6, 20, 35 and 76 cm. Water collected at depths 70 and 76 cm is assumed to reflect
142 the amount and quality of water leaving the root zone and will be used in the calculation of leaching. Field
143 observations (Elberling & Ladegaard-Petersen, 2005) suggested that root densities below 70-80 cm in both
144 stands were very low, supporting the assumption that lysimeter installed at this depth sampled water lost
145 from the root zone. A suction of -35 kPa was used to extract soil solution at 20 second intervals with a
146 period of 60 seconds in between extractions. The sampling containers were placed inside a wooden box,
147 buried in the forest floor in order to limit the suction required to extract water, avoid freezing and suppress
148 biological activity. Water samples were placed in a dark room at $5 \text{ }^\circ\text{C}$ shortly after extraction and sub-
149 samples for analysis were taken within 24 hours of sampling. Sampling of the accumulated amounts of soil

150 water followed the time schedule of TF with the exception that sampling of soil solution in the N. spruce
151 stand terminated in July 2000.

152

153 2.5. Laboratory analyses

154 Soil pH was measured in distilled water (1:2.5). All other chemical soil analyses were made using only
155 the soil fraction finer than 2 mm. Total organic carbon (TOC) was measured after acidification using the dry
156 combustion method at 1250 °C on an Eltra SC-500 analyzer, with an accuracy of $\pm 0.2\%$. Total N was
157 analyzed using dry combustion and infrared detection of N using a LECO FP-428, version 2.03 apparatus.
158 The grain size distribution was analyzed after samples were oxidized with 4M H₂O₂ to remove organic
159 matter. After drying, samples were sieved through meshes of 63, 125, 250, 500, 1000 and 2000 μm . The
160 fraction finer than 20 μm was analyzed using a hydrometer (Gee & Bauder, 1986). Soil water retention
161 curves were obtained using a pressure membrane apparatus at pressures equivalent to 10, 100, 1000 and
162 15,000 cm of water corresponding to pF 1.0, pF 2.0, pF 3.0 and pF 4.2 (Klute, 1986).

163 Conductivity and pH of the recovered water was determined upon return to the laboratory the same day.
164 This was also the case for the alkalinity, which was determined by titration of HCl. The rest of the water was
165 kept at 5°C and dark until analyzed. Total dissolved Fe, Mn, Mg, Ca, K, and Na were determined on
166 acidified water samples using atomic absorption spectroscopy and Cl, NO₃ and SO₄ were determined on
167 water samples of non-preserved water using ion chromatography.

168

169 2.6. Meteorological variables

170 Measurements of meteorological data were collected in a nearby forest, Stenholt Vang, located 2 km from
171 the Nødebo site. At Stenholt Vang a 10 m mast was installed in a clearing and measurements were made at
172 15 minute intervals (described by Hansen, 2003) and used as input values in the modelling for both stands.
173 The meteorological variables include: precipitation at 2 m wind speed at 10 m, relative humidity at 2 m,
174 global radiation at 2 m in and air temperature at 2 m. In this study the meteorological variables have been
175 modified to represent daily means of wind speed, relative humidity, global radiation and temperature.

176 Precipitation is represented as daily accumulation in mm. Fig. 2C shows the temporal variation of
177 precipitation (mm) from 1999 – 2001 measured at Stenholt Vang (see Hansen, 2003). Annual observed
178 precipitation in 2000 was 798 mm. The maximum input of daily precipitation (53 mm) occurred on
179 September 2. The pattern of precipitation showed no distinct trend during the year, but daily inputs of
180 precipitation exceeding 30 mm all occurred from June – September.

181

182 2.7. Model description and setup

183 The CoupModel is a one-dimensional numerical model that takes the vegetation, soil and atmosphere into
184 account. Evapotranspiration forms a central part of the model governing the input of water to the soil.
185 Evapotranspiration can be divided into three parts: evaporation from the soil surface, evaporation of
186 intercepted water in the canopy and transpiration from the plants. The actual evapotranspiration is calculated
187 as the sum of evaporation from intercepted water, soil evaporation and transpiration. The forest canopy is
188 represented by a single leaf concept as given by Monteith (1965), for calculation of both direct evaporation
189 losses from intercepted water and transpiration from the leaf originating from the water uptake from the soil
190 (Jansson & Karlberg, 2004). The actual transpiration is calculated on the basis of the potential transpiration
191 given by the Penman-Monteith formulation and response functions for soil and meteorological factors
192 (Jansson & Karlberg, 2004). Soil evaporation is considered by using an energy balance approach (Alvenäs &
193 Jansson, 1997). When modelling water balance of forest ecosystems, key input parameters include: LAI,
194 surface resistance of canopy, vertical root distribution and soil hydraulic properties such as unsaturated and
195 saturated hydraulic conductivity (Jansson *et al.*, 1999b).

196 Input variables included air temperature, wind speed, global radiation, relative humidity and precipitation.
197 The flow of water in the soil is calculated on the basis of Richard's equation using an explicit numerical
198 solution using finite differences either with a forward or a central difference scheme (Jansson & Karlberg,
199 2004). In the CoupModel the soil was divided into eight layers, 0 – 0.2, 0.2 – 0.4, 0.4 – 0.6, 0.6 – 0.8, 0.8 –
200 1.2, 1.2 – 1.6, 1.6 – 2.0 and 2.0 – 2.5 m for each stand. The grain size distribution and retention curve
201 observed for each of the horizons provided the basis for estimating the hydraulic properties. Hydraulic

202 properties (lambda [shape parameter of the water retention curve], air entry, residual water, wilting point,
203 tortuosity, matrix and total conductivity) were calculated in the CoupModel using the Brooks-Corey
204 formulation for the retention curve and the Mualem formulation for the hydraulic conductivity (Jansson &
205 Karlberg, 2004). Retention curves and texture analysis were made only on samples from one pit in each
206 stand. Thus, it was not possible to determine statistical differences between stands. As retention
207 measurements at pF 4 failed for the N. spruce stand and due to similarities in textural properties for the two
208 stands, soil characteristics for the beech stand were used for both stands. LAI was used to estimate the
209 interception capacity of canopy precipitation and also the partitioning of the global short wave radiation
210 between canopy and soil surface. The vertical root distribution defines the zone from which water uptake
211 occurs and therefore the amount of water available for transpiration.

212 Adjustment of surface resistance, soil physical properties (lambda and tortuosity), water capacity per LAI
213 and temperature coefficients controlling water uptake by plants were based on observed values of the
214 volumetric water content of the soil. All parameter values used to adjust the CoupModel are given in table 2a
215 and table 2b. Model performance was evaluated on the basis of the coefficient of determination for a linear
216 regression between simulated and observed values (R^2), root mean square error (RMSE) and the mean error
217 (ME). Statistical results of the model simulations are shown in Table 3.

218 The simulation runs from July 10, 1998 to August 9, 2001 with daily output values. The investigated
219 period was the year 2000. Measurements of mean water content in 0 – 20, 0 – 40, 0 – 60, 0 – 80 and 0 – 120
220 cm through the entire period were used to fit the model to observed data. Measurements of meteorological
221 variables (see section 2.6) were assumed to be similar for both stands and used as input accordingly.
222 Information about the vegetation was taken both from the Nødebo site (tree height, root distribution) and a
223 location in Jutland named Ulborg (Hansen, 2003). The tree heights were set to 9 m in the beech stand and 23
224 m in the N. spruce stand. Absolute value of canopy resistance has been modified according to tree species
225 and annual variations simulated as suggested by Person & Lindroth (1994). Water uptake was defined as a
226 pressure head approach, where water uptake is calculated on the basis of response functions for water content
227 and soil temperature (Jansson & Karlberg, 2004). The start of the growing season (and the corresponding
228 water uptake) was defined with a triggering temperature approach (Jansson & Karlberg, 2004). The growing

229 season began when the day length exceeded 10 hours and the accumulated temperature was above 9 °C. It
230 ended when the day length became less than 10 hours. As the beginning and end of the growing season is
231 determined on the basis of meteorological variables the length of the growing season is identical for the N.
232 spruce and beech.

233 Water leaving the lower boundary was used as a measure of the percolation from the forest ecosystem.
234 Outputs of percolation are used to estimate leaching of elemental fluxes from January – December 2000.
235

236 2.8. Calculation of elemental fluxes in stemflow, throughfall and soil water

237 The annual input of the elements (Ca, Mg, K, Na, Fe, Al and Cl, SO₄-S, NO₃-N) in TF in kg ha⁻¹ y⁻¹ was
238 calculated by multiplying the concentration of elements (mg L⁻¹) by the amount of water collected in the
239 funnel converted to mm ha⁻¹. For the input of elements from SF, it was assumed that the tree from which SF
240 was collected was representative of the entire stand. The number of trees ha⁻¹ was multiplied by the amount
241 of water collected (in L) and afterwards multiplied by the concentration (mg L⁻¹) and converted to kg ha⁻¹
242 y⁻¹. Seasonal trends and total leaching of the elements were calculated using model output of percolation
243 (mm day⁻¹). The soil water was continuously extracted and sampled roughly on a weekly basis. Observed
244 element concentrations in extracted soil water were assumed to represent the mean concentrations during the
245 extraction time. Subsequently, daily values of percolation were multiplied by element concentrations to
246 calculate daily values of leaching and finally converted to monthly and annual values.

247

248 2.9 Statistical analyses

249 Statistical analyses applied in this paper included simple linear regression calculating
250 Pearson's coefficient of explanation, R² (Eq. 7), on observed pairs of values (Jansson & Karlberg,
251 2004). Significance was tested using a 95% confidence level, and relationships were significant if R²_{obs} >
252 R²_{crit95%} implying the p ≤ 0.05. R²_{crit95%} were looked up in a table containing critical Pearson's coefficients
253 of explanation. Furthermore, mean error (ME) and root mean square error (RMSE) used in this
254 paper were calculated using the CoupModel.

255 Mean error was calculated using the following equation:

$$256 \frac{\sum sim(i) - \sum obs(i)}{n}$$

257 where $sim(i)$ and $obs(i)$ are the values at the i 'te observation and n is the number of observations.

258 Root mean square error was calculated according to the following equation:

$$259 \sqrt{\frac{\sum (sim(i) - obs(i))^2}{n}}$$

260 where $sim(i)$ and $obs(i)$ are the values at the i 'te observation and n is the number of observations.

261

262

263 **3. Results**

264 3.1. Observed and simulated water content

265 Time series of measured soil water storage (SWS) in mm for beech and N. spruce are shown in Figs. 2A
266 and B. The temporal variation of SWS is similar for both stands and shows a distinct trend with the highest
267 values during winter and spring, consistent with precipitation events, declining during May and reaching
268 minimum values through the summer and early autumn. For the upper two layers the level of soil water (16 –
269 133 mm) storage is equal in both stands. The difference between the stands increases with depth. For the
270 entire soil profile (0 – 120 cm) the level of SWS for beech is between 137 – 320 mm and 107 – 272 mm for
271 N. spruce. The simulated SWS is shown as solid lines and reveals an acceptable fit to observations (Fig. 2
272 and Table 3) and within the error bars for replicate measurements observed for the 0 – 120 cm layer. The
273 coefficient of determination for a linear regression (R^2) between observed and simulated water contents in
274 the entire soil profile (0-120 cm) is 0.97 ($p < 0.001$) for beech with a ME of –2.7 mm, equalling 1% of the
275 mean simulated SWS. For N. spruce R^2 is 0.91 ($p < 0.001$) with a ME of – 3.5 mm (2% of mean simulated
276 SWS).

277

278 3.2. Water balance simulations

279 Simulated yearly outputs (Table 4) and monthly values of precipitation (P), actual transpiration (Et),
280 actual interception evaporation (Ei), actual soil evaporation (Es) and percolation (A) (Figs. 2A, B and C)
281 reveal important differences between the two tree species. Fig. 3A shows P on a monthly basis in the year
282 2000, which shows that there is no tendency in the variation of P during the year. Fig. 3B shows the monthly
283 water balance for beech. Actual evapotranspiration (Ea), the sum of Et, Ei and Es, shows a clear temporal
284 variation with the highest values (71 – 112 mm) from May – September. Actual evapotranspiration exceeds
285 P from May – August, except in June. For beech, Et and Ei is low (0 – 2 mm) from January – April, thus Es
286 dominates evaporation from February – April. From May – July, Et increases (34.5 – 86.4 mm) and
287 gradually decreases (67.3 – 3.5 mm) from August – November and becomes very low in December (0.1
288 mm), whereas Ei remains relatively constant (14.9 – 32 mm) from May – November. The annual share of Ea
289 in relation to P is 68% and equals 581 mm. Interception loss contributes with 18% of P equalling 158 mm on
290 an annual basis. The annual Et is 339 mm, which constitutes 40% of P. Annual Es is four times lower than
291 Ea and contributes annually with 10% of P (equal to 84 mm). In relation to Ea the shares of Ei, Et and Es are
292 27%, 58% and 14% respectively. The annual percolation from the beech stand is 292 mm and constitutes
293 34% of P. From February – April, the percolation is at a maximum (54 – 80 mm per month), it declines from
294 May – August (33 – 6 mm) and reaches a minimum from September – December of 4 – 2 mm per month.
295 Table 4 shows that throughfall (TF) equals soil surface infiltration (SI), thus indicating that surface runoff is
296 unlikely and is consistent with lack of surface runoff as observed in the field.

297 Fig. 3C shows the water balance for the N. spruce stand. Generally, the temporal variation of Ea is much
298 like that for beech except that Ea values are several times higher in January – March than what was
299 calculated for the beech stand. Ea reaches the maximum from May – September (81 – 136 mm). The
300 minimum Ea occurs in January at 28 mm. As opposed to the beech stand, Ei remains high throughout the
301 year (23 – 43 mm). Transpiration in the N. spruce stand generally shows the same temporal variation as in
302 the beech stand and values are higher in most cases. From April – July, Et increases from 12.3 – 80 mm, it
303 remains constant through August – September (63 – 67 mm), declines to 15 mm in October, and reaches the

304 minimum in November and December (0.2 – 3.1 mm). Es is constantly low and varies little (0.2 – 6.2 mm)
305 throughout the year. The annual Ea is 823 mm and constitutes 96% of P. The annual interception loss is 396
306 mm, which equals 46% of P. The annual Et is 388 mm (45% of P). Soil evaporation amounts to 39 mm and
307 constitutes only 5% of P on an annual basis. The division of Ea into the shares of Ei, Et and Es shows that Ei
308 and Ea contribute equally with 48% and 47% of Ea, and Es constitutes 5%. Annual TF (464 mm) equals soil
309 infiltration (SI).

310 As the input of water is the same in N. spruce and beech it can be deduced that the percolation in the N.
311 spruce stand is lower. The annual percolation from spruce is 41 mm, which constitutes only 5% of P. The
312 temporal variation of percolation on a monthly basis is shown in Fig. 3C. The percolation is at a maximum
313 from May – July (6 – 8 mm) and is low (1 – 3 mm) from January – April and October – December. It is seen
314 that the maximum of percolation is displaced in both stands compared to the minimum of Ea at the beginning
315 of the year. This is due to the fact that the percolation is a measure of the water flow at 2.5 m depth and thus
316 delayed compared to inputs at the surface.

317

318 3.3. Element concentrations and fluxes in throughfall, stemflow and soil water

319 Table 5 shows the annual mean concentrations (mg L^{-1}) and fluxes ($\text{kg ha}^{-1} \text{y}^{-1}$) of Ca, Mg, K, Na, Fe, Al
320 and Cl, $\text{SO}_4\text{-S}$, $\text{NO}_3\text{-N}$ in TF, SF and soil water below the root zone (see section 2.4) for the beech and N.
321 spruce stands. Concentrations of elements in TF from the spruce stand generally exceed those in the beech
322 plot. This is reflected in the fluxes of TF as all elements, except Al, show the largest flux in the N. spruce
323 stand. Sodium and Cl fluxes between the two stands from TF are notable, as the input of Cl and Na is three
324 times higher in the N. spruce stand. As it can be concluded from Table 5, the flux from SF is less than 10%
325 of the flux from TF in both stands for all elements. The flux from SF of Mg, K, Al and $\text{NO}_3\text{-N}$ is largest in
326 beech. The amount of leaching of the different elements is generally largest in the N. spruce stand, but the
327 leaching of Ca^{2+} , Fe and $\text{NO}_3\text{-N}$ from beech exceeds that from the spruce. The most conspicuous differences
328 in leaching between the two stands are seen for the following elements: $\text{NO}_3\text{-N}$ (beech: 39, N. spruce: 0.5),
329 Ca (beech: 65, N. spruce: 6), Na (beech: 16, N. spruce: 19), Cl (beech: 39, N. spruce: 47). The difference in
330 leaching of $\text{NO}_3\text{-N}$ is especially notable, because the mean annual concentration of $\text{NO}_3\text{-N}$ (11.3 mg L^{-1}) in

331 the beech stand equals the maximum limit of NO₃-N for drinking water in Denmark (Ministry of
332 Environment, 1988). The consistently high concentrations of NO₃-N in the soil water, as indicated by the low
333 standard deviation of 4.3, cause the high annual leaching of 39 kg ha⁻¹ y⁻¹, while it is very small in the N.
334 spruce stand at only 0.5 kg N ha⁻¹ y⁻¹ which is reflected in the low soil water concentration. The trend and
335 magnitude of monthly leaching of NO₃-N from the beech and N. spruce stands are shown in Fig. 4. It can be
336 seen that the leaching from the spruce stand during the whole period is several orders of magnitude smaller
337 than the corresponding values for the beech stand, and the peak of NO₃-N leaching is displaced towards the
338 summer for N. spruce.

339

340 **4. Discussion**

341 4.1. Simulations of water balance

342 Fitting of the model showed that it was possible to simulate water percolation based on measurements of
343 volumetric water content converted to water storage in mm. The statistically significant ($R^2 = 0.91 - 0.97$,
344 $p < 0.001$) simulations are supported by the low ME of -2.7 and -3.5 for beech and N. spruce respectively.
345 Based on the fitting of the CoupModel it is assumed that the respective water balances are representative of
346 the two forest ecosystems at Nødebo. The same conclusion at different locations was made by Ladekarl
347 (2001) showing similar patterns of percolation from the forest soils. The values of percolation can therefore
348 be used for both stands to calculate the leaching of elements (Fig. 3B and 2C).

349 A comparison of the water balance in beech and N. spruce reveals several distinct differences. In Table 4
350 the main constituents of the water balance are related to P. Actual evapotranspiration is largest in N. spruce,
351 exceeding Ea for the beech stand by 29%. If the shares of Ei, Et and Es in relation to P are compared, the
352 difference in Ea between N. spruce and beech is mainly due to differences in Ei. Transpiration is highest in
353 N. spruce (45% of P) but is the same order of magnitude as in beech (40% of P). In both cases, Ei is low (5
354 and 10%). For Ei the shares are 18% for beech and 46% for N. spruce, thus the interception loss in spruce is
355 more than twice of that in beech. A high interception loss in spruce forests was also reported by Alavi et al.
356 (2001) and Mossin & Ladekarl (2003). Mossin & Ladekarl also concluded that a high interception loss

357 would lead to low percolation. Ladekarl (2001) compared the water balances at several locations in Great
358 Britain, Germany, France and Denmark and concluded that there was no significant difference in Et between
359 beech and N. spruce. For the investigations listed in Ladekarl (2001) Et was in the range of 255 – 398 mm
360 for beech forests and 204 – 400 mm for N. spruce forests, and the values simulated for the beech stand at
361 Nødebo (Table 4) are in the same order as these values, but it is seen that the N. spruce stand at Nødebo is
362 much lower than reported values. Ladekarl (2001) concluded that the main differences in the water balance
363 between beech and N. spruce located on the same soil type were due to the differences in interception loss,
364 which is also the case in this study.

365 Difference in interception was related to LAI and a stem/branch related component for the beech stand.
366 Phenological observations at Nødebo (Elberling & Ladegaard-Petersen, 2005) show that the beech trees set
367 leaf at the end of April/beginning of May and defoliate during October, which is similar to observations
368 made at Ulborg. It is therefore assumed that the temporal variation of interception loss in beech and N.
369 spruce at Nødebo is simulated satisfactorily.

370 The absolute level of interception is determined by the parameterisation of the model. A high interception
371 will reduce TF and A. The amount of TF can indicate whether the interception is estimated correctly. The
372 amounts of TF measured at Nødebo are 450 mm for beech and 300 mm for N. spruce. The simulated TF is
373 698 mm and 464 mm for beech and spruce, respectively. As the TF samplers used at Nødebo were open at
374 the top, evaporation from the samplers is expected to reduce the collected amounts. It is not likely that
375 evaporation from the TF funnels accounts for the entire difference between measured and simulated values.
376 Because only two replicates of TF were installed in each stand it is possible that the true variation in TF
377 amounts is not represented in the collected amounts at Nødebo.

378 Water balance simulations made for Stenholt Vang in the period of 1995 – 1997 (Bastrup-Birk *et al.*,
379 2003) showed that percolation constituted between 36 and 30% of P for beech and between 26 and 22% for
380 N. spruce. Percolation in the beech stand at Nødebo fell within these values, but was lower in the spruce
381 stand. Both in terms of absolute amount, as compared to the nearby location of Stenholt Vang, and the
382 relative difference between beech and N. spruce stands at different locations in Europe, the water balance for
383 Nødebo was determined satisfactorily.

384 4.2. Element leaching

385 In order to validate the water balance, observed concentrations of Cl^- were used as a conservative
386 element. The input and output fluxes of Cl^- exceeded the amount of any other investigated element (Table 5)
387 and the concentration of Cl^- was higher in SW as compared to TF (Table 5). Stemflow was not considered
388 further as the fluxes of elements were low as compared to TF (Table 5). Thus, the increase in concentration
389 of Cl^- in the soil water could be used to calculate the loss of water due to evaporation. The ratios of $\text{Cl}_{\text{TF}}/\text{Cl}_{\text{SW}}$
390 for beech and N. spruce were 0.49 and 0.31, respectively, showing that the amount of water input at the
391 forest floor had been reduced by 51% and 69% through transpiration and soil evaporation. Using the model
392 results (Table 4) the corresponding evaporation of TF from Et and Es was 61% and 92% for beech and N.
393 spruce, respectively. The Cl^- approach underestimated the evaporation, and the differences in evaporation
394 between beech and N. spruce using Cl^- concentration underestimated the calculated differences between the
395 two stands. An explanation for this could be, that observed concentrations in TF on an annual basis cannot be
396 compared to concentrations of Cl in soil solution as the input of Cl in TF is delayed compared to output from
397 the root zone. Despite the inability of the Cl^- approach to validate calculated evapotranspiration it still
398 showed that evaporation was much higher in the N. spruce stand. Subsequently, the CoupModel has been
399 used to predict the tree-specific leaching of $\text{NO}_3\text{-N}$ as a function of the water balance. As previously shown
400 by Kennedy & Pitman (2004), NO_3 concentrations below the root zone in British soils have been
401 successfully explained by differences in soil water contents and water balance.

402 If the same approach is used in the case of Nødebo, the concentration of $\text{NO}_3\text{-N}$ in the water leaving the
403 root zone can be calculated as a function of input (TF) and evaporation (Cl^-). Thus, an average concentration
404 of 1.3 mg L^{-1} $\text{NO}_3\text{-N}$ in TF results in a 2-fold increase in concentrations. However, the actual observed
405 concentrations are roughly 10 times higher (Table 5), which indicates that other inputs than TF are
406 responsible for the actual concentrations. This is in contrast to Kennedy & Pitman (2004). A $\text{NO}_3\text{-N}$
407 enrichment is not seen for the N. spruce stand, which indicates that part of the added N from TF is taken up
408 during downward transport.

409 The total input of elements at Nødebo is generally larger in N. spruce than in beech (Table 5), indicating
410 that the atmospheric deposition is largest in spruce. This was also concluded by Rothe et al. (2002). The

411 investigation published in Rothe et al. (2002) encompassed several locations across Europe and it was stated
412 that the higher canopy deposition in N. spruce stands caused the leaching of NO₃-N to be highest in spruce
413 compared to beech stands. Kristensen et al. (2004) showed that soil solution N was higher in conifers than in
414 broadleaves when throughfall input of N was below 10 kg ha⁻¹ y⁻¹. These findings are in contrast to the
415 results obtained at Nødebo even though the input of N in N. spruce follows the trend presented in Rothe et al.
416 (2002).

417

418 4.3. Forest soil N dynamics

419 The leaching of NO₃-N in beech is much higher than the leaching from N. spruce. This could only partly
420 be explained by the lower percolation from the spruce stand as compared to the beech stand. Thus,
421 differences in NO₃-N leaching from the two stands must include an analysis of the soil N dynamics in the
422 two forest soils.

423 The work by Callesen et al. (1999) classified Danish forests soils on the basis of NO₃-N concentrations in
424 soil solution. The investigation showed that >60% of the forests had an annual leaching below 2 kg N ha⁻¹ y⁻¹
425 (median concentration < 2 mg N L⁻¹) and had a low risk of leaching of NO₃-N below the root zone. Seven
426 percent of the investigated forests had a median concentration of NO₃-N exceeding drinking water standards
427 at 11.3 mg N L⁻¹ and had a high risk of leaching of NO₃-N below the root zone. If the two stands at Nødebo
428 are compared to the findings in Callesen et al. (1999), the N. spruce stand has a low risk of NO₃-N leaching
429 whereas the beech stand has a large risk of NO₃-N leaching below the root zone. The high leaching of NO₃-N
430 from the beech stand suggests that input of atmospheric N in the beech stand exceeds the rate of uptake and
431 the forest ecosystem could possibly be saturated with nitrogen as defined in Aber et al. (1989) and
432 Gundersen (1991). The low concentrations of NO₃-N in the lowest horizon in the N. spruce stand could
433 indicate that N added from TF and SF and N released from decomposition is taken up by vegetation and
434 immobilized by microorganisms during the downward transport in the soil. Gundersen et al. (2006) stated
435 that thinning of the stand only affected the NO₃-N leaching to a minor degree, and as the latest thinning at
436 Nødebo occurred in 1998, it is estimated that the effect of the thinning in 1998 was diminished at the start of
437 the measurement period in December 1999.

438 Gundersen et al. (1998a) discussed the possibility using the C:N ratio of the forest floor to indicate the
439 degree of NO₃-N leaching and classified the forests on the basis of the C:N ratio, with: C:N > 30 as N-
440 limited and low risk of leaching; 25 < C:N < 30 as intermediate and moderate risk of leaching; C:N < 25 as
441 N-saturated and high risk of leaching. Dise et al. (1998), Borken & Matzner (2004), Kristensen et al. (2004)
442 support the findings in Gundersen et al. (1998a). Absolute amounts of leaching based on the C:N ratio were
443 not proposed due to the great variance between locations with the same C:N ratio. The relationship between
444 C:N of the forest floor and NO₃-N leaching was most evident for coniferous species but the use of C:N for
445 deciduous species needs further investigation.

446 The low level of leaching agrees with the C:N of the forest floor in the N. spruce stand (~36) and is thus
447 characterized as N-limited. Excess N is probably assimilated by the microorganisms and transferred to
448 stable pools, humus or aboveground biomass. The occurrence of NO₃-N below the root zone in the N.
449 spruce stand reveals that the N-cycling is not completely tight and it may be expected that transport of N
450 occurs by convective mass transfer with percolating water in larger pores, which agrees with the findings in
451 section 4.2.

452 By contrast, the C:N ratio for the upper horizon in the beech stand is ~20 and is characterized as N
453 saturated. The variation of the C:N ratio with depth shows a relative stable C:N in the A – B2ws horizons (0
454 – 40 cm) and a minimum of 8 in the B3ws (37 – 57 cm), see Table 1. According to Gundersen et al. (1998a)
455 the beech stand can be characterized as N-saturated, which agrees with the high amount of N leached. The
456 use of C:N has mainly found application in coniferous forests, but works well at Nødebo for both types of
457 forest.

458 The reason for the differences in the C:N ratio of the upper horizon is to be found in the properties of the
459 organic substrate. At Nødebo there is no profound accumulation of an organic horizon in the beech stand but
460 a pronounced O-horizon in the N. spruce plot. This indicates that the soil fauna decompose all newly added
461 litter over one year in the beech stand, probably because the need for N in the microbial community has been
462 satisfied (Johnson, 1992). This pattern is typical of nutrient-rich soils (Callesen et al., 1999). Due to the
463 properties of N. spruce needles and the associated resistance to microbial decomposition, input of mineral N

464 through TF and SF has longer mean residence time. Consequently, a relatively smaller amount of N are
465 released into the soil water and consistent with an accumulation in the O-horizon.

466 Johnson (1992) reviewed N retention in forest soil and argued that non-biological N retention was small
467 in acidic soils. As the soil $\text{pH}_{\text{H}_2\text{O}}$ in both stands at Nødebo varies between 3.7 – 4.9, non-biological retention
468 of N is expected to be small. Therefore, the microbial community is expected to cause the N retention in both
469 stands at Nødebo. The leaching of $\text{NO}_3\text{-N}$ in the beech suggests that the microbial community has a different
470 composition with more nitrifying organisms responsible for release of $\text{NO}_3\text{-N}$ into the soil solution (Zhong &
471 Makeschin, 2004) than it is the case of the N. spruce stand. A constant addition of N from the atmosphere
472 will increase the amount of nitrifying organisms and thus increase the soil solution concentration of NO_3^-
473 (Johnson, 1992; Zhong & Makeschin, 2004). In turn, the nitrification is inhibited in the N. spruce stand due
474 to the low concentration of $\text{NO}_3\text{-N}$ in the soil water.

475 Because both stands are situated relatively close to the edge of the forest, an edge-effect (Beier &
476 Gundersen, 1989) could play a part in the elevated concentrations of $\text{NO}_3\text{-N}$ below the root zone.
477 Spangenberg & Kölling (2004) found elevated fluxes of ions in TF at the forest edge and leaching. An edge-
478 effect is only possible when winds are from an eastern direction because the two forest stands are situated in
479 the easternmost part of the forest. It can be seen from Fig. 5 that a change toward eastern winds occurs in the
480 spring and autumn. A deposition of N in the spring when leaves are absent in the beech stand could lead to
481 increased amounts of N deposited directly on the forest floor compared to when the trees have leaves. In
482 combination with low microbial activity, low vegetative uptake and high rates of water percolation in the
483 winter and spring, the deposited N could leach into the soil and be transported unattended with the soil water
484 beneath the root zone. Therefore, it is indicated that a certain edge-effect exists at Nødebo, at least for the
485 beech stand. It was shown in section 4.2. that only a fraction of the $\text{NO}_3\text{-N}$ could be explained by input from
486 TF. It is not possible to identify a trend between the amount of input of N and leaching of N from the C-
487 horizon. Gundersen *et al.* (1998b) found that $\text{NO}_3\text{-N}$ leaching was correlated with an N-status of the
488 ecosystem and not significantly correlated with N-deposition in coniferous and deciduous forest ecosystems,
489 which supports the findings in this paper.

490

491 **5. Conclusion**

492 This study has shown that the CoupModel, a process oriented SVAT-model, was useful to document
493 differences in water balance between two contrasting forest stands located in close proximity to each other
494 on similar soil types. Using water balance modelling, chemical analysis of soil water and few geochemical
495 parameters, a broad view of ecosystem functioning has been established, with respect to both the
496 geochemical cycling of N and other important nutrients and the mass balance of water. It is concluded that
497 the model outputs can be used to calculate monthly as well as annual fluxes of leaching from the root zone.

498 The main difference in the water balance was caused by interception loss through evaporation from the
499 canopy, as the transpiration was in the same order of magnitude in the two species. The annual percolation
500 was 292 mm from the beech stand and 41 mm from the N. spruce stand.

501 Leaching of elements in $\text{kg ha}^{-1} \text{y}^{-1}$ was largest from the N. spruce stand and is probably due to the higher
502 canopy deposition in spruce trees. The leaching of $\text{NO}_3\text{-N}$ differed considerably between the two species, as
503 $\text{NO}_3\text{-N}$ leaching from the beech stand was $39 \text{ kg NO}_3\text{-N ha}^{-1} \text{y}^{-1}$, compared to $0.5 \text{ kg ha}^{-1} \text{y}^{-1}$ in the N. spruce
504 stand. On the basis of the leaching of $\text{NO}_3\text{-N}$ it was concluded that the beech stand was possibly saturated
505 with respect to N, due to excess atmospheric input of N in relation to the N uptake of the trees. By contrast,
506 the N. spruce stand could be characterized as unsaturated with respect to N. The difference in $\text{NO}_3\text{-N}$
507 leaching between the two species could be explained by several factors. The rich nutrient status of the soil in
508 combination with a C:N ratio between 8 – 20 in the beech stand soil suggests that the need for N in the
509 microbial community and vegetation is fulfilled through litter and soil organic matter decomposition and that
510 the demand for external sources of N is small. Furthermore, a possible edge-effect could cause increased
511 inputs of atmospheric N to the forest floor in the spring in the beech stand, leading to excess input of N in
512 relation to demand. The small leaching of $\text{NO}_3\text{-N}$ from the N. spruce stand suggests a high need for N in the
513 vegetation and microbial community, thus increasing the retention of N. The N status of the two forest
514 ecosystems shows that structure and functionality of the microbial community is different for the two stands,
515 leading to differences in N retention and N leaching. The functionality and response of the two forest

516 ecosystems in relation to the water balance and atmospheric deposition of N observed at Nødebo suggests
517 that further attention is needed when selecting tree species for future afforestation projects.

518

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622

624 **Table 1. Main soil characteristics at Nødebo for beech and Norway spruce, including carbon content, C:N ratio,**
625 **pH(H₂O), bulk density (ρ) and weight percentages of clay, silt and sand.**

Beech						Norway spruce					
Horizons	Depth (cm)	C (mg g ⁻¹)	C:N	pH (H ₂ O)		Horizons	Depth (cm)	C (mg g ⁻¹)	C:N	pH(H ₂ O)	
A1	0 – 10	53	20.8	3.7		O	8 – 0	376	36.5	3.7	
E(B)	10 – 20	26	22.6	4.0		E	0 – 11	36	32.7	3.8	
B2WS	20 – 37	23	22.4	4.5		B2	11 – 29	32	31.9	4.2	
B3WS	37 – 57	5.3	8.0	4.8		B3	29 – 43	8.7	21.9	4.4	
C1	57 – 100	0.74	11.5	4.9		C1	43 – 85	2.4	-	4.5	
C2	100 – 150	1.5	30.1	-		C2	85 –	2.8	-	4.4	
Horizons	Depth (cm)	ρ (g cm ⁻³)	clay (<2) %	silt (2-63) %	sand (63-2000) %	Horizons	Depth (cm)	ρ (g cm ⁻³)	clay (<2) %	silt (2-63) %	sand (63-2000) %
A1	0 – 10	0.84	6.2	25.9	67.9	O	8 – 0	-	-	-	-
E(B)	10 – 20	1.2	10.8	22	67.2	E	0 – 11	1.2	8.2	25.1	66.7
B2WS	20 – 37	1.1	11.5	22.6	65.9	B2	11 – 29	1.1	8.2	25.4	66.4
B3WS	37 – 57	1.3	10	17.2	72.8	B3	29 – 43	1.2	6.3	23.1	70.6
C1	57 – 100	1.5	9.5	17	73.5	C1	43 – 85	1.5	8.3	25.5	66.2
C2	100 – 150	1.3	9	18.3	72.7	C2	85 –	1.6	14.1	46.7	39.2

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Table 2a. Parameter values used to adjust the CoupModel. Model parameters assigned to default values are not included.

Variable	Name	Unit	Beech	Norway spruce
Water Capacity Base independent of LAI	WaterCapacityBase	mm	1	0
Water Capacity per LAI	WaterCapacityPerLAI	mm m ⁻²	0.5	0.5
Within Canopy Resistance	WithinCanopyRes	s m ⁻¹	0.5	0.5
Altitude of meteorological station	AltMetStation	m	50	50
Altitude of site	AltSimPosition	m	50	50
Reference height above displacement height of respective stand	ReferenceHeight	m	10	10
Rate coefficient for surface runoff from soil surface pool	SurfCoef	-	0.1	0.1
Maximum surface pool without generation of surface runoff	SurfPoolMax	mm	10	10
Minimum soil hydraulic conductivity	MinimumCondValue	mm d ⁻¹	1E-4	1E-4
Latitude	Latitude	-	56	56
Critical threshold for water uptake	CritThresholdDry	cm water	1000	1000
Power coefficient for sensitivity of water uptake to potential transpiration rate	NonDemandRelCoef	-	0	0
Aggregate sorption coefficient in matrix domain	AScaleSorption	-	1	1
LAI	LAI	m ² m ⁻²	0 – 4.5	8 – 8.5
Canopy surface resistance	Resistance surface	s m ⁻¹	50 – 500	40 – 500

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Table 2b. Soil physical properties used in the CoupModel for the beech and Norway spruce stands at Nødebo. Soil physical properties were calculated on the basis of retention analysis. Lambda represents a shape parameter of the water retention curve.

Depth (m)	Lambda (-)	Air entry (cm)	Saturation (%)	Wilting point (%)	Residual water (%)	Matrix cond. (mm d ⁻¹)	Total cond. (mm d _i)	Tortuosity (-)
0 – 0.1	0.195	5.3	55	6.5	0.1	3870	3870	1
0.1 – 0.2	0.188	1.5	57	6.5	0.1	5715	5715	1
0.2 – 0.37	0.186	1.0	57	6.5	0.1	4712	4712	1
0.37 – 0.57	0.228	4.0	44	2.5	0.1	3000	3000	1
0.57 – 1	0.235	2.1	42	2.5	0.09	3000	3000	1
1 – 1.5	0.266	1.2	40	2.5	0.06	3000	3000	1

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635 **Table 3. Statistical performance of the CoupModel.**

Stand	Horizon (cm)	R ²	RMSE (mm)	ME (mm)	Mean measured (mm)	n
<u>Norway spruce</u>	0 – 20	0.82	7.0	1.1	48.6	60
	0 – 40	0.86	16.2	13.2	77.7	60
	0 – 60	0.91	13.9	6.1	116.1	60
	0 – 80	0.87	21.3	16.1	132.1	60
	0 – 120	0.91	16.7	-3.5	199.6	60
Beech	0 – 20	0.74	7.0	4.8	48.3	63
	0 – 40	0.68	22.2	19.0	79.7	63
	0 – 60	0.85	20.2	15.7	120.1	63
	0 – 80	0.93	28.0	26.5	140.1	63
	0 – 120	0.97	10.3	-2.7	231.0	63

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Table 4. Annual simulated output in mm beech and N. spruce using the CoupModel. Outputs are also given as percentages of precipitation.

	Beech (mm)	% of P	Norway spruce (mm)	% of P
<u>Precipitation</u>	856	-	856	-
<u>Transpiration</u>	339	40	388	45
<u>Interception loss</u>	158	18	396	46
<u>Soil evaporation</u>	84	10	39	5
<u>Evapotranspiration</u>	581	68	823	96
<u>Soil infiltration</u>	692	-	461	-
<u>Throughfall</u>	698	82	464	54
<u>Percolation</u>	292	34	41	5

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644 **Table 5. Annual mean concentrations in mg L⁻¹ of throughfall (TF), stemflow (SF) and soil water (SW) below the**
645 **root zone for the beech and N. spruce stands. Standard deviation for the different elements is given in**
646 **parenthesis. TF flux, SF flux and leaching are the corresponding fluxes in kg ha⁻¹ y⁻¹. Letters indicate whether**
647 **the values for the Norway spruce stand are greater than (a), equal to (b) or less than (c) the corresponding value**
648 **for the beech stand.**

Beech						
	TF (mg L ⁻¹)	TF flux (kg ha ⁻¹ y ⁻¹)	SF (mg L ⁻¹)	SF flux (kg ha ⁻¹ y ⁻¹)	SW (mg L ⁻¹)	Leaching (kg ha ⁻¹ y ⁻¹)
Ca	2.3 (1.5)	10.7	2.6 (1.3)	0.23	20.7 (5.0)	64.6
Mg	0.8 (0.4)	3.5	1.3 (1.4)	0.14	1.9 (0.4)	6.3
K	1.8 (3.4)	9.5	3.0 (1.8)	0.27	0.3 (0.2)	0.9
Na	3.4 (3.5)	14.2	1.8 (8.0)	0.98	5.0 (0.9)	15.6
Fe	0.02 (0.03)	0.06	0.01 (0.02)	0.0009	0.04 (0.04)	0.1
Al	0.2 (0.3)	1.04	0.2 (0.3)	0.02	1.4 (0.4)	4.6
Cl	5.8 (6.0)	26.8	13.7 (23.2)	1.67	11.9 (4.8)	38.7
SO ₄ -S	1.2 (0.5)	5.7	2.0 (1.3)	0.20	4.4 (1.1)	11.4
NO ₃ -N	1.3 (0.8)	5.2	1.4 (0.9)	0.14	11.3 (4.3)	39.1
N. spruce						
Ca	4.7 (2.2) (a)	14.3 (a)	12.4 (11.2) (a)	0.29 (a)	10.1 (2.3) (c)	6.4 (c)
Mg	3.9 (4.6) (a)	9.34 (a)	4.7 (6.3) (a)	0.12 (c)	9.0 (3.6) (a)	4.7 (c)
K	5.5 (5.0) (a)	15.4 (a)	10.5 (7.6) (a)	0.24 (c)	2.2 (1.0) (a)	1.5 (a)
Na	11.4 (11.5) (a)	42.5 (a)	33.0 (21.4) (a)	1.01 (a)	49.5 (22.8) (a)	18.3 (a)
Fe	0.1 (0.3) (a)	0.46 (a)	0.1 (0.1) (a)	0.002 (a)	0.1 (0.2) (a)	0.02 (c)
Al	0.2 (0.4) (b)	0.94 (c)	0.3 (0.5) (a)	0.01 (c)	11.6 (13.6) (a)	2.3 (c)
Cl	25.7 (21.2) (a)	88.8 (a)	69.9 (70.3) (a)	1.96 (a)	82.1 (47.0) (a)	46.7 (a)
SO ₄ -S	3.6 (1.3) (a)	10.2 (a)	8.0 (4.8) (a)	0.21 (a)	22.8 (0.6) (a)	9 (c)
NO ₃ -N	3.0 (2.0) (a)	6.62 (a)	2.0 (1.9) (a)	0.04 (c)	0.6 (0.7) (c)	0.5 (c)

649 Figure texts

650 **Fig. 1** The Nødebo study site. The letters indicate the different stands. A: beech, B: Norway spruce
651 and C: Norway spruce damaged during a storm in December 1999. In this study stands A and B
652 were investigated.

653

654 **Fig. 2.** Simulations (solid lines) versus measurements of water storage (mm) for beech (A) and
655 Norway spruce (B) in 1999 – 2001. The following layers are represented: 0 – 20 (■), 0 – 40 (▲), 0
656 – 60 (△), 0 – 80 (○) and 0 – 120 (●) cm. Error bars are shown for the 0 – 120 layer.

657

658 **Fig. 3.** Monthly values of precipitation (P) and water balance elements (evaporation and deep
659 percolation) in mm for beech (A) and Norway spruce (B) stands at Nødebo in 2000. In A and B
660 evaporation is shown as positive values and divided into transpiration (Et), interception evaporation
661 (Ei) and soil evaporation (Es). The deep percolation (A) is represented as negative values.

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663 **Fig. 4.** Monthly values of leaching (kg ha^{-1}) of $\text{NO}_3\text{-N}$ from beech (◇) and Norway spruce (■) in
664 2000 at Nødebo.

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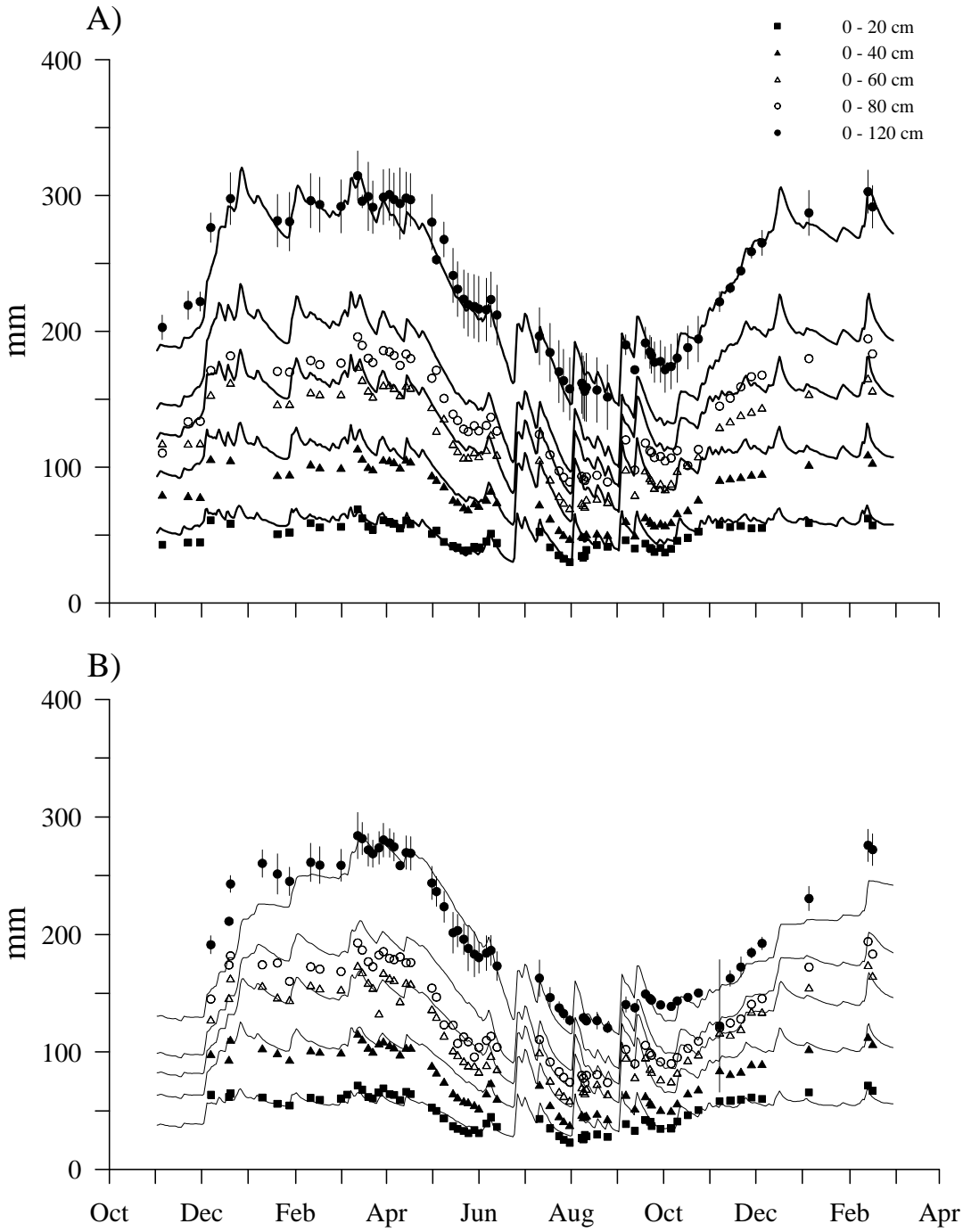
666 **Fig. 5.** Wind direction from the meteorological station (Hansen (ed.), 2003). Values are floating
667 mean values of 500 measurements. Records of wind direction were stored every 10 minutes. Values
668 of the y-axis are designated with letters representing eight directions, with a 45° increment between
669 values.



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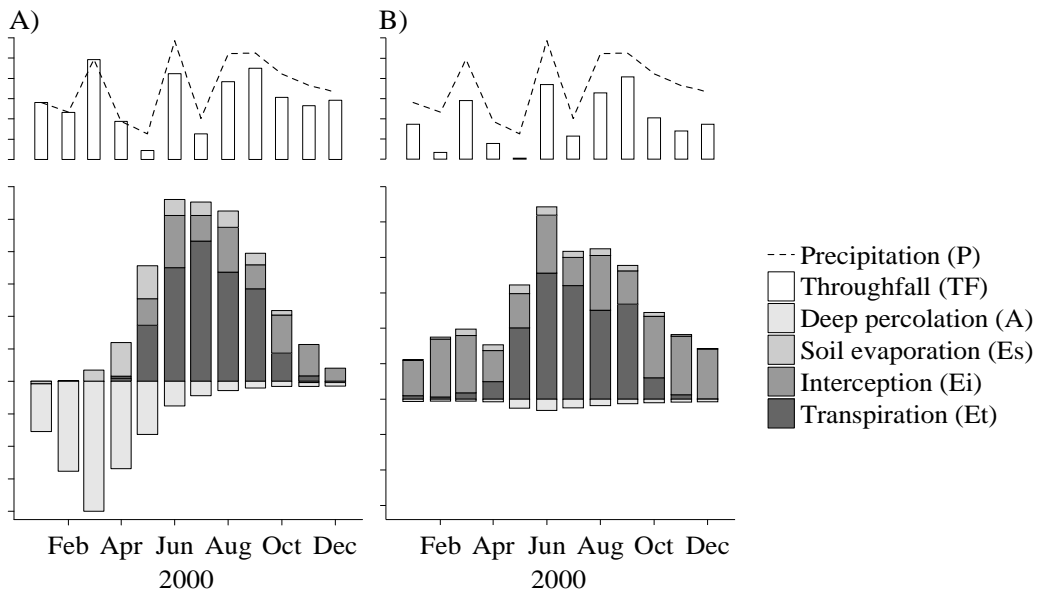
Fig. 1

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Fig. 2

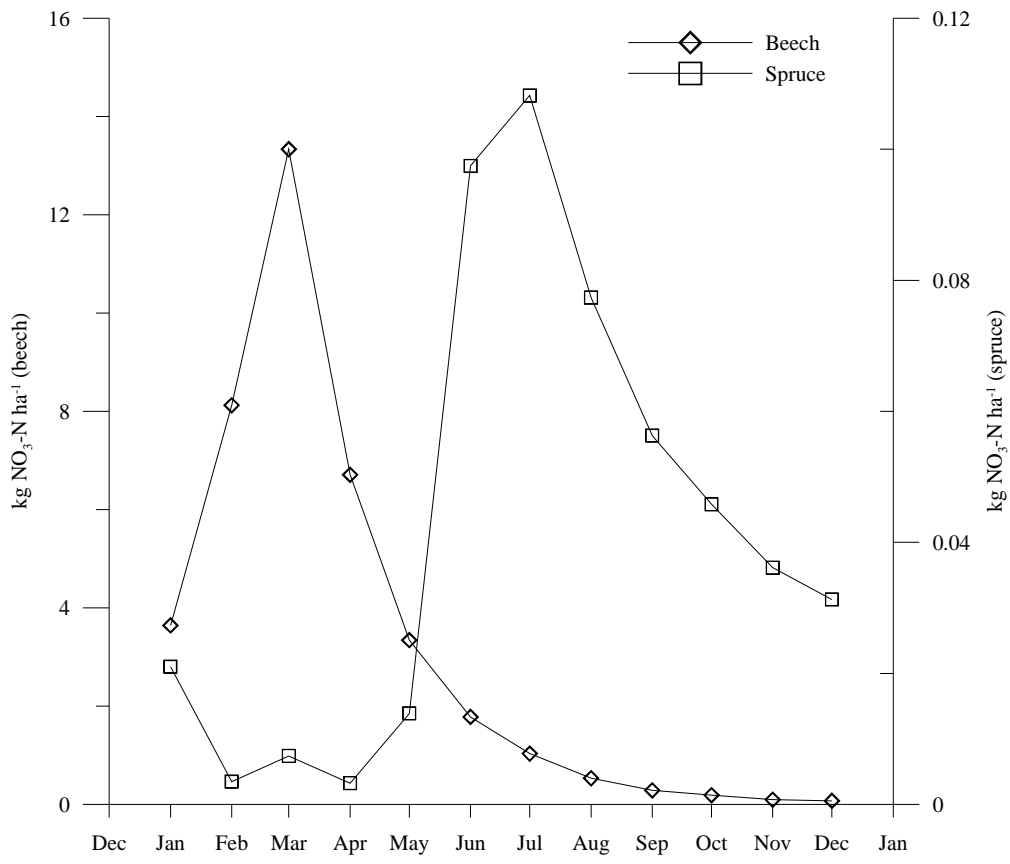
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Fig. 3

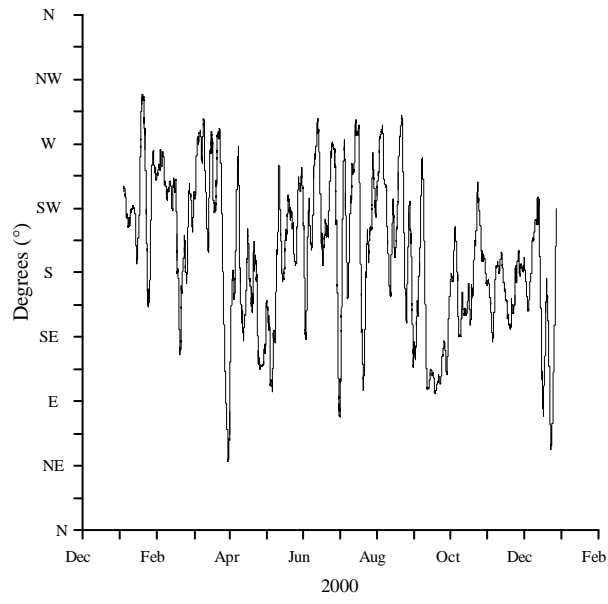


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Fig. 4



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Fig. 5.